



Two ways to account for the dependence of microwave observations of precipitation on uncertain microphysics



Z.S.Haddad (JPL), J. Steward (UCLA), S. Kacimi (JPL), J. Munchak (GSFC),
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Explicitly account for habit and size distribution assumptions, making sure that **additional simplifying assumptions are consistent** with observations

or

Try to **smooth over habit and size assumptions** to make sure that analysis does not try to adjust poorly modeled parameters (and that it adjusts prognostic variables consistently with covariances)



⇒ Assume closed-form diameter distributions (e.g. Γ):

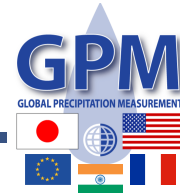
$$N(D) = N_0 D^\mu e^{-\Lambda D}$$

Let's try to interpret the parameters in terms of physically meaningful quantities:

$$D_m = \frac{\int D D^3 N(D) dD}{\int D^3 N(D) dD} = \frac{\mu + 4}{\Lambda}$$

$$q = \int \frac{4}{3} \pi \left(\frac{D}{2} \right)^3 \rho N(D) dD = \frac{\pi \rho \Gamma(\mu + 1) D_m^{\mu+1}}{6 (\mu + 4)^{\mu+1}} N_0$$

$$\sigma_m = \left(\frac{\int (D - D_m)^2 D^3 N(D) dD}{\int D^3 N(D) dD} \right)^{1/2} = \frac{D_m}{\sqrt{\mu + 4}}$$



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⇒ Assume closed-form diameter distributions (e.g. Γ):

In the models, typically assume N_0 constant and $\mu = 0$.

What that implies is:

$$\begin{aligned} q &= \int \frac{4}{3} \pi \left(\frac{D}{2} \right)^3 \rho N(D) dD = \frac{\pi \rho \Gamma(\mu + 1) D_m^{\mu+1}}{6 (\mu + 4)^{\mu+1}} N_0 \\ &= \frac{\pi \rho}{24} N_0 D_m \end{aligned}$$



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What that implies is:

$$D_m = \frac{24}{\pi \rho N_0} q$$

In particular,

- $D_m/q = \text{constant}$, and
- $\max(D_m)/\min(D_m) = \max(q)/\min(q)$



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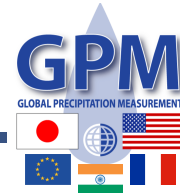
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But $3.5 \text{ mm} / 0.5 \text{ mm} \neq \max(R^{0.9})/\min(R^{0.9}) \approx 100 \text{ mm/hr} / 0.1 \text{ mm/hr}$



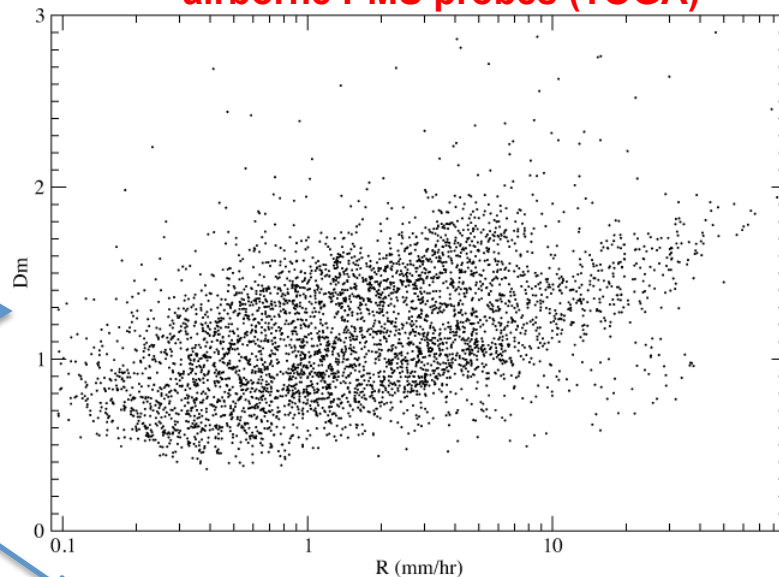
⇒ In fact, hydrometeor data suggest $D_m \sim q^{0.2} \pm \text{white noise}$

$D_m \sim q^{0.2}$ behavior in profiler data

is consistent with TOGA-COARE

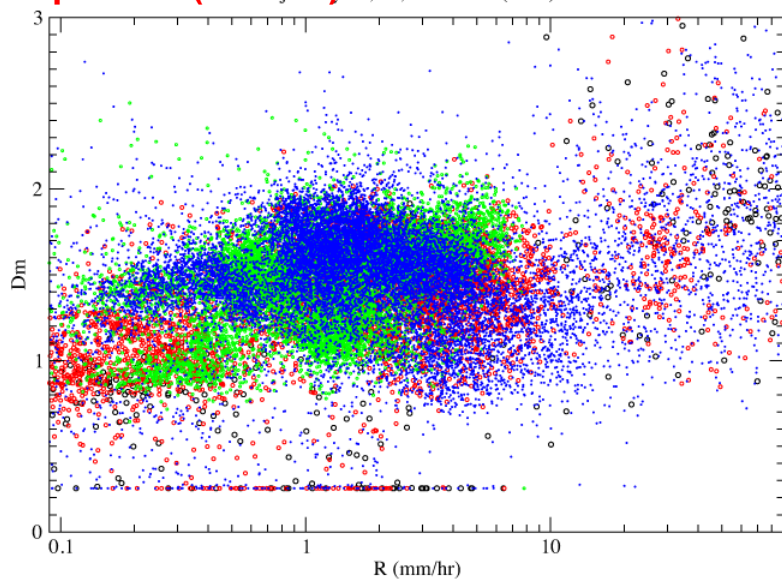
and Kwajex data ...

airborne PMS probes (TOGA)

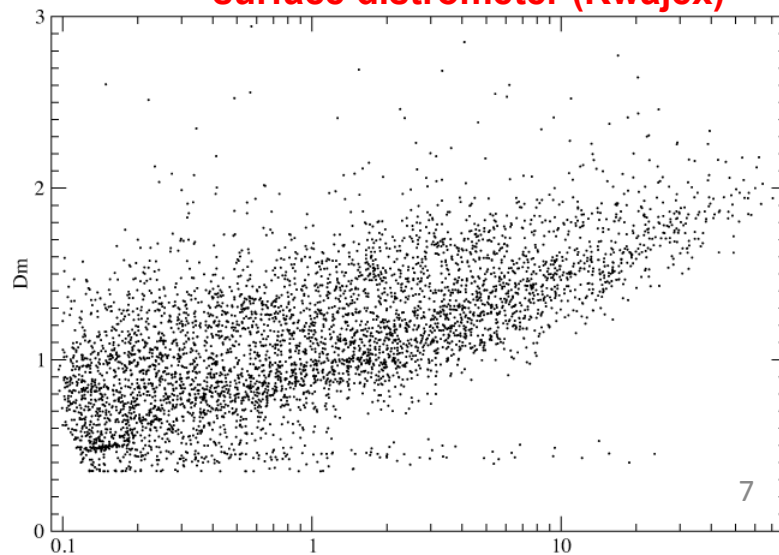


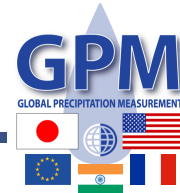
Doppler profiler (Darwin)

January 19, 20, 22 and 23 (2006)



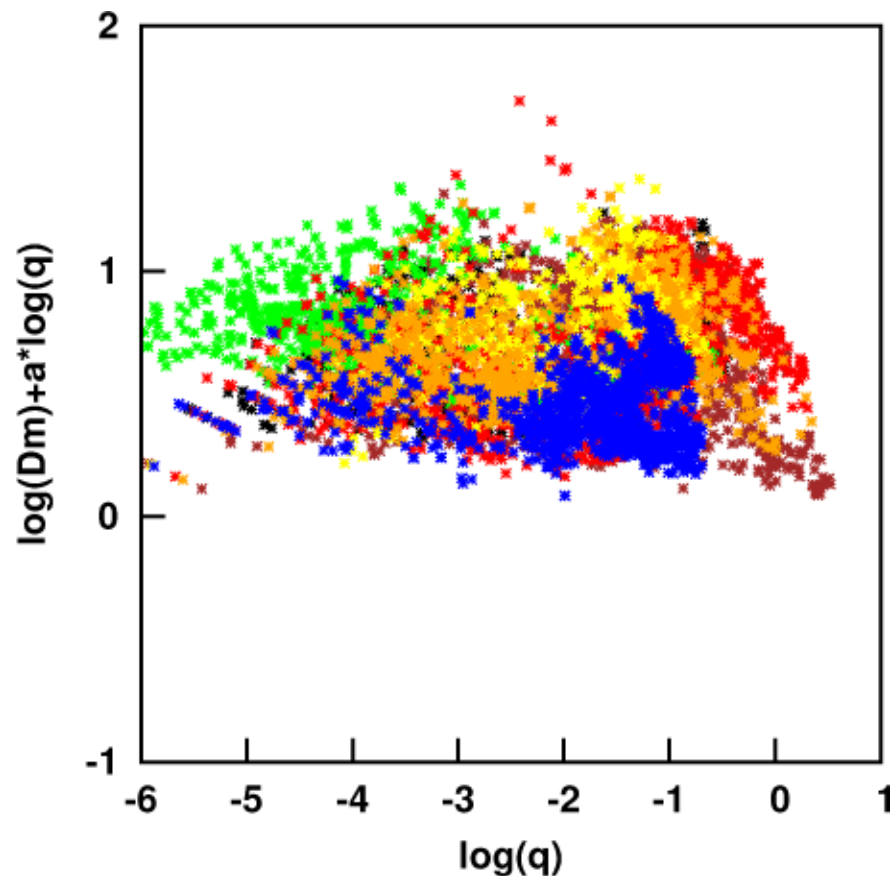
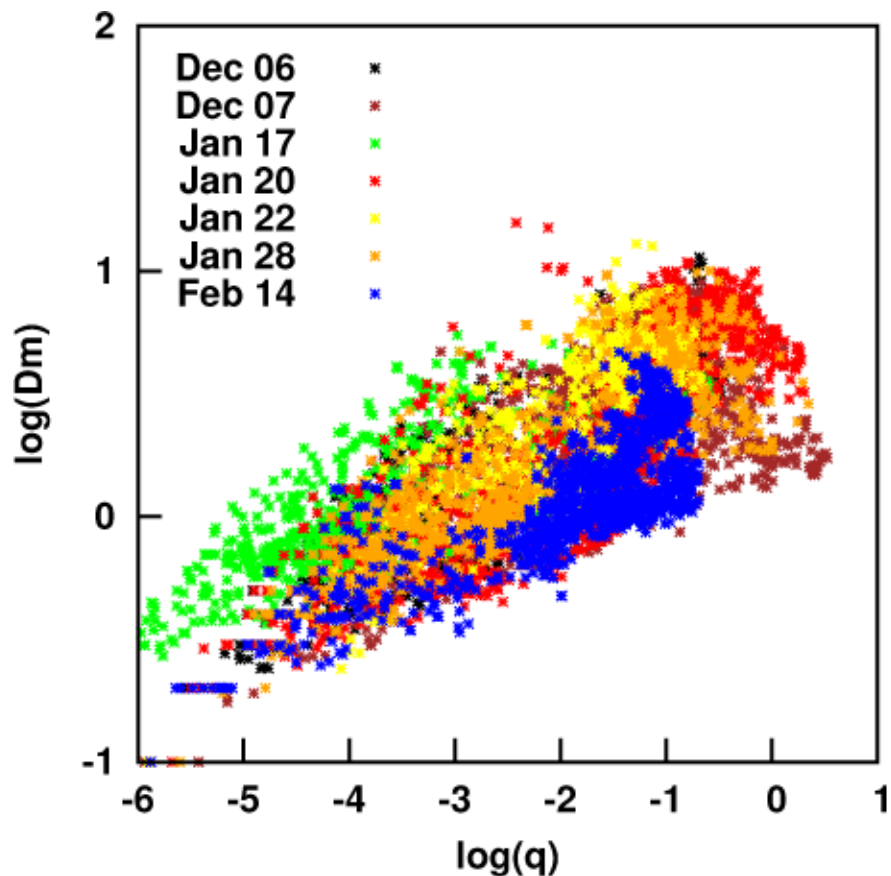
surface distrometer (Kwajex)





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2D video Distrometer (King City, ON – thanks to Environment Canada)





⇒ Assume closed-form diameter distributions (e.g. Γ):

$$N(D) = N_0 D^\mu e^{-\Lambda D}$$

Fixing Λ is at least as problematic:

$$D_m = \frac{\int D D^3 N(D) dD}{\int D^3 N(D) dD} = \frac{\mu + 4}{\Lambda}$$

$$\sigma_m = \left(\frac{\int (D - D_m)^2 D^3 N(D) dD}{\int D^3 N(D) dD} \right)^{1/2} = \frac{D_m}{\sqrt{\mu + 4}}$$

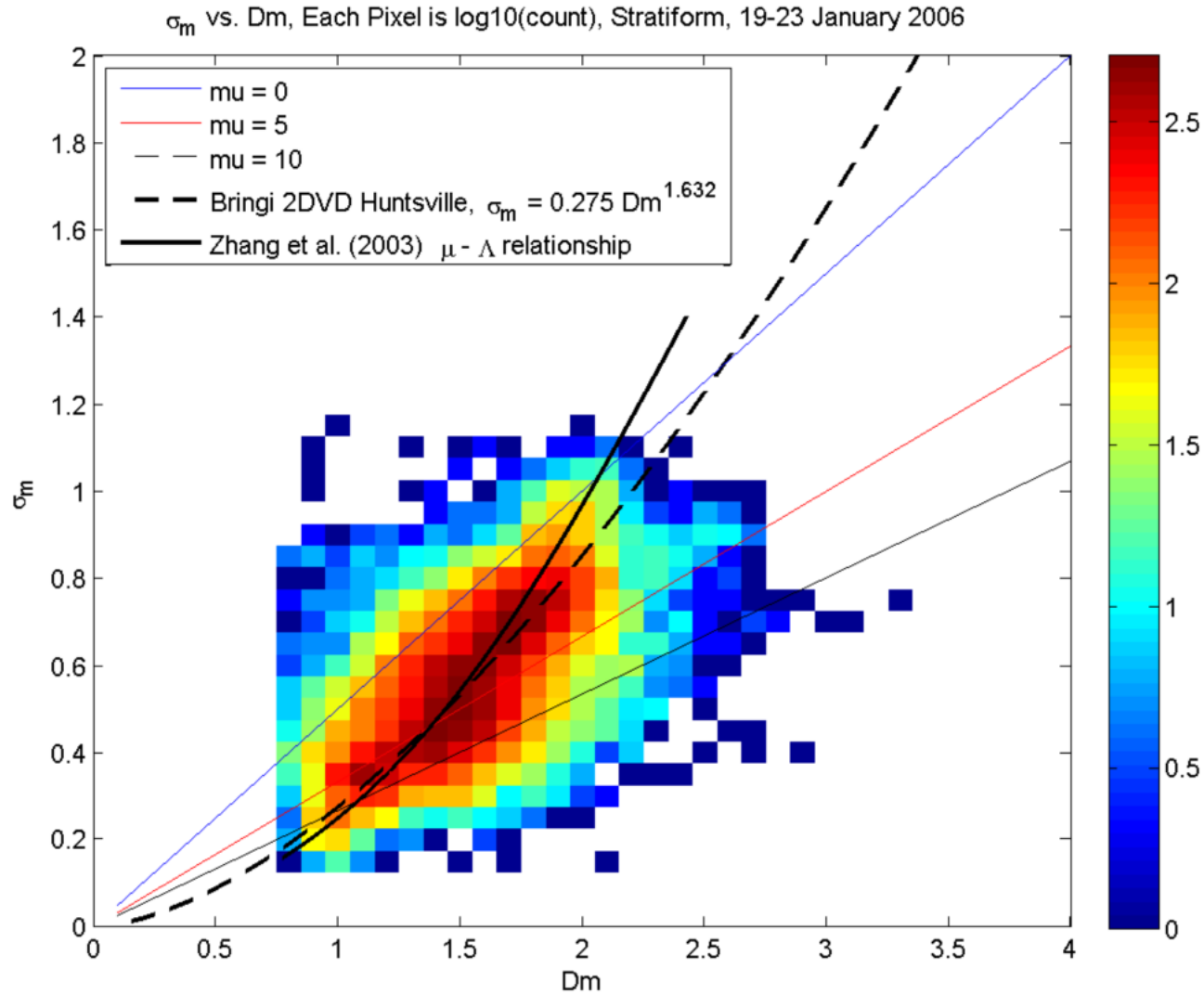
The above imply:

$$D_m = \frac{D_m^2 / \sigma_m^2}{\Lambda} \Rightarrow \frac{D_m}{\sigma_m^2} = \text{const} \quad , \text{ i.e. } \quad \sigma_m = \text{const} D_m^{0.5}$$



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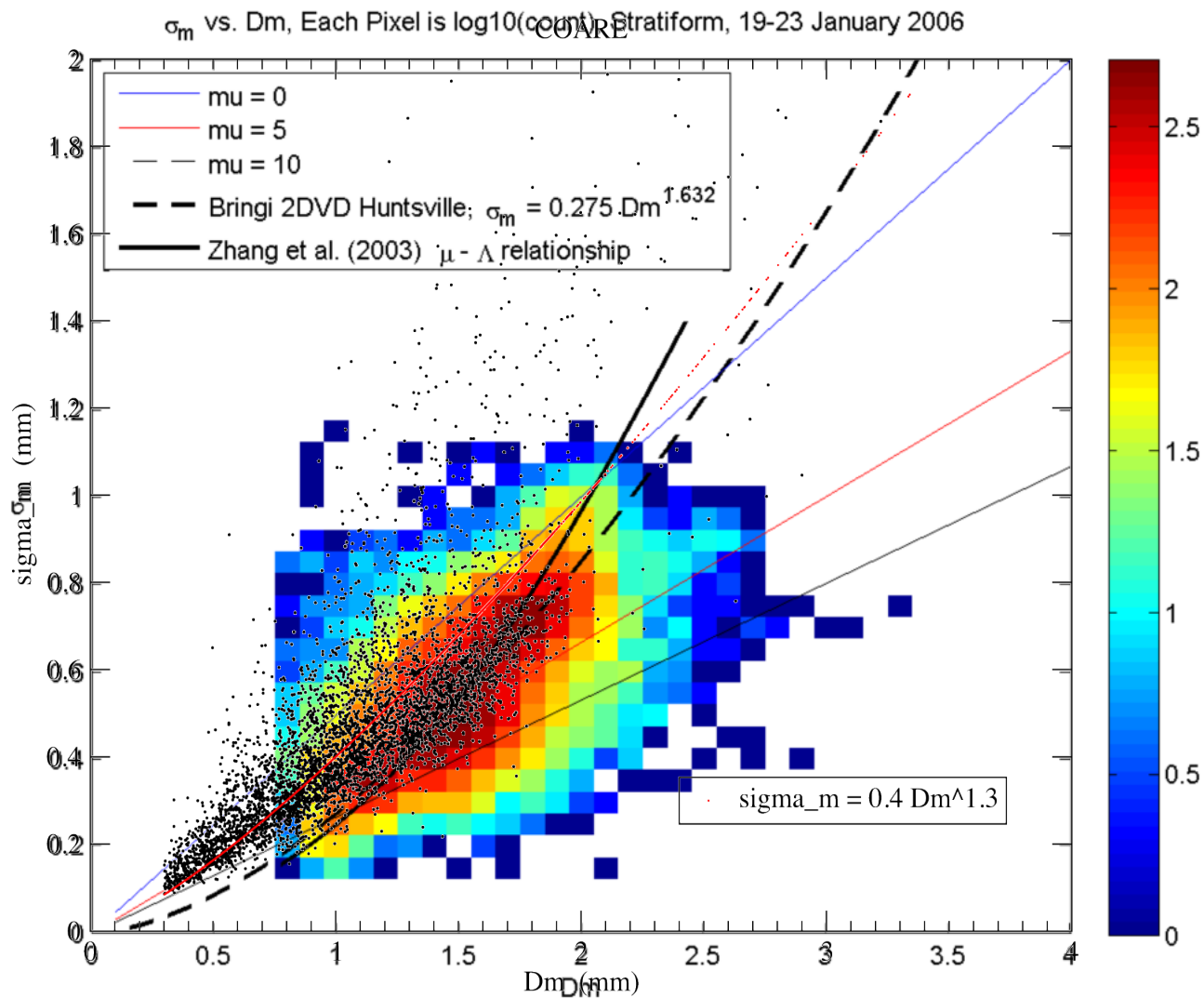
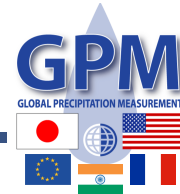


$$\sigma_m \neq \text{const } D_m^{0.5}_{10}$$



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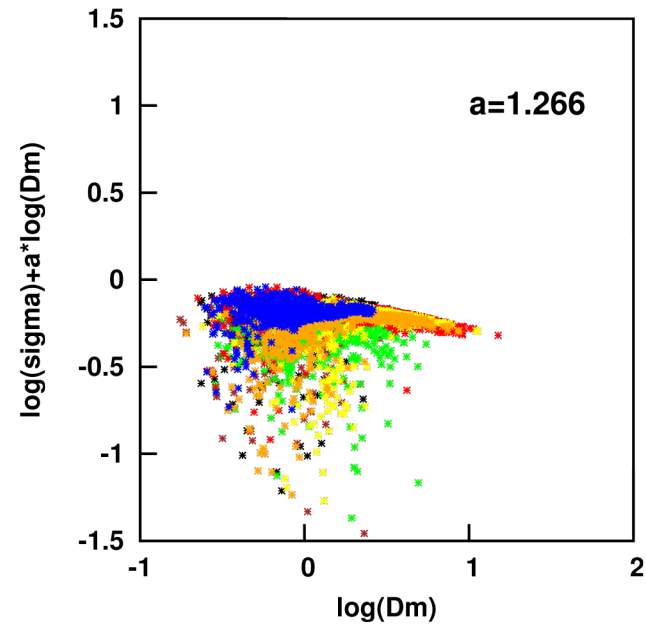
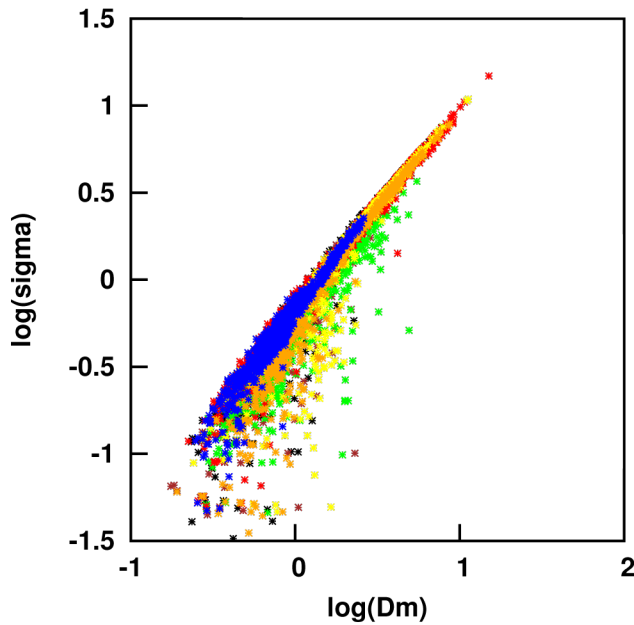


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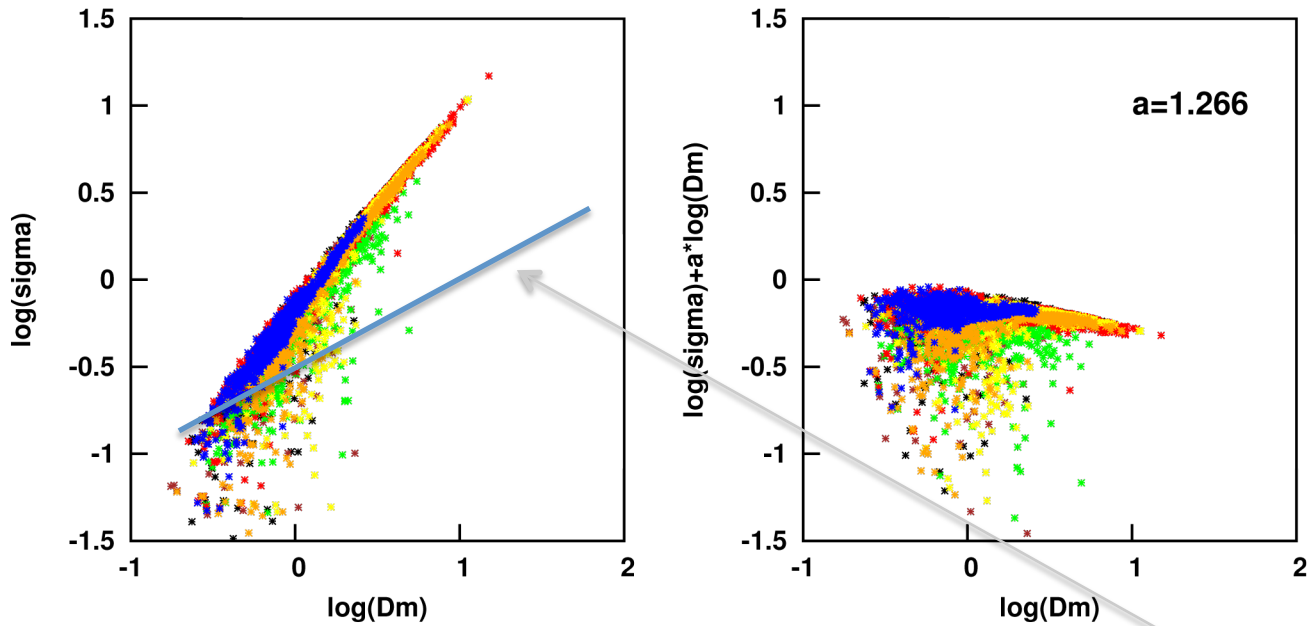
2D video Distrometer (King City, ON – thanks to Environment Canada)



$$\sigma_m \neq \text{const } D_m^{0.5}$$

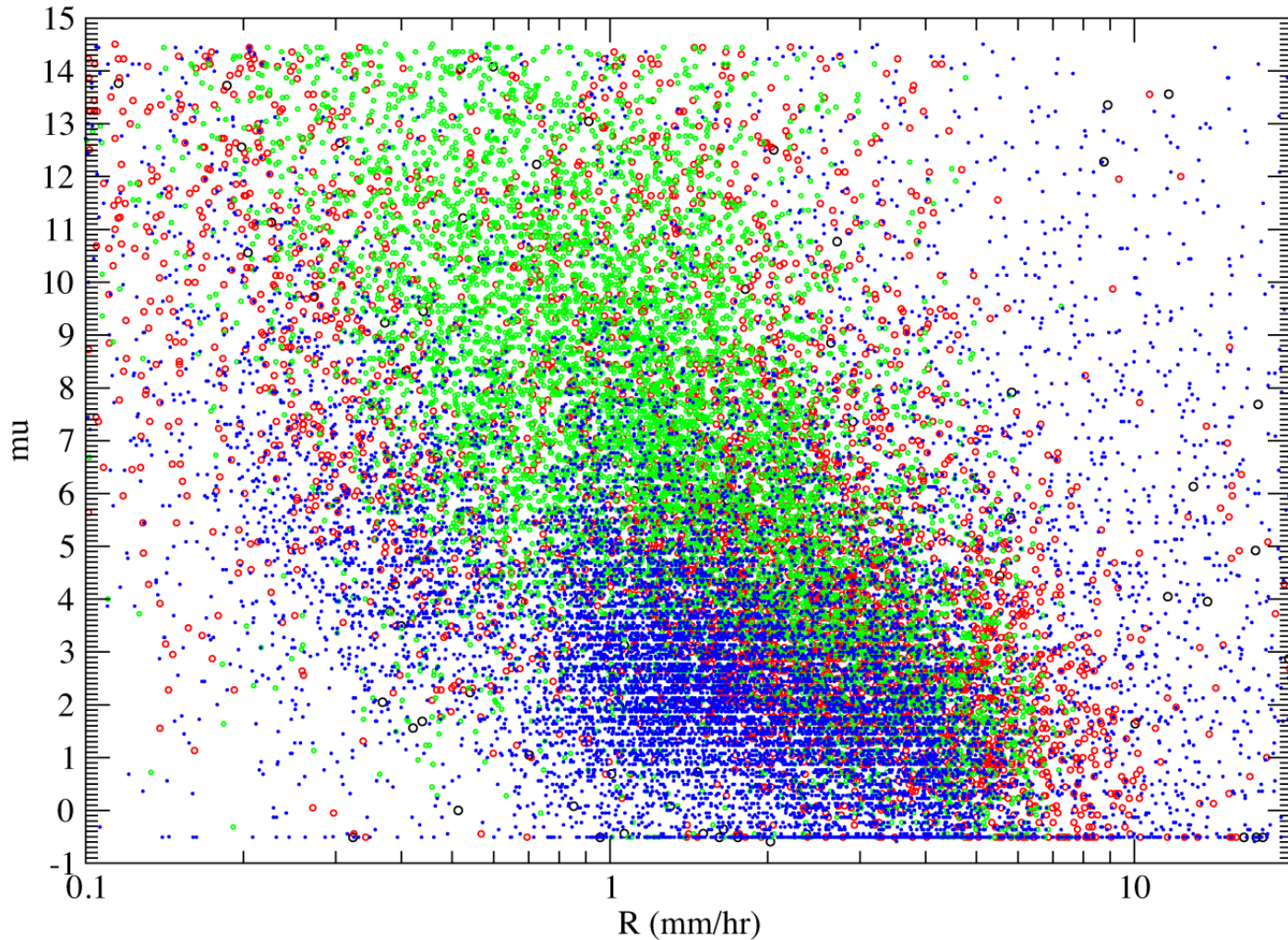


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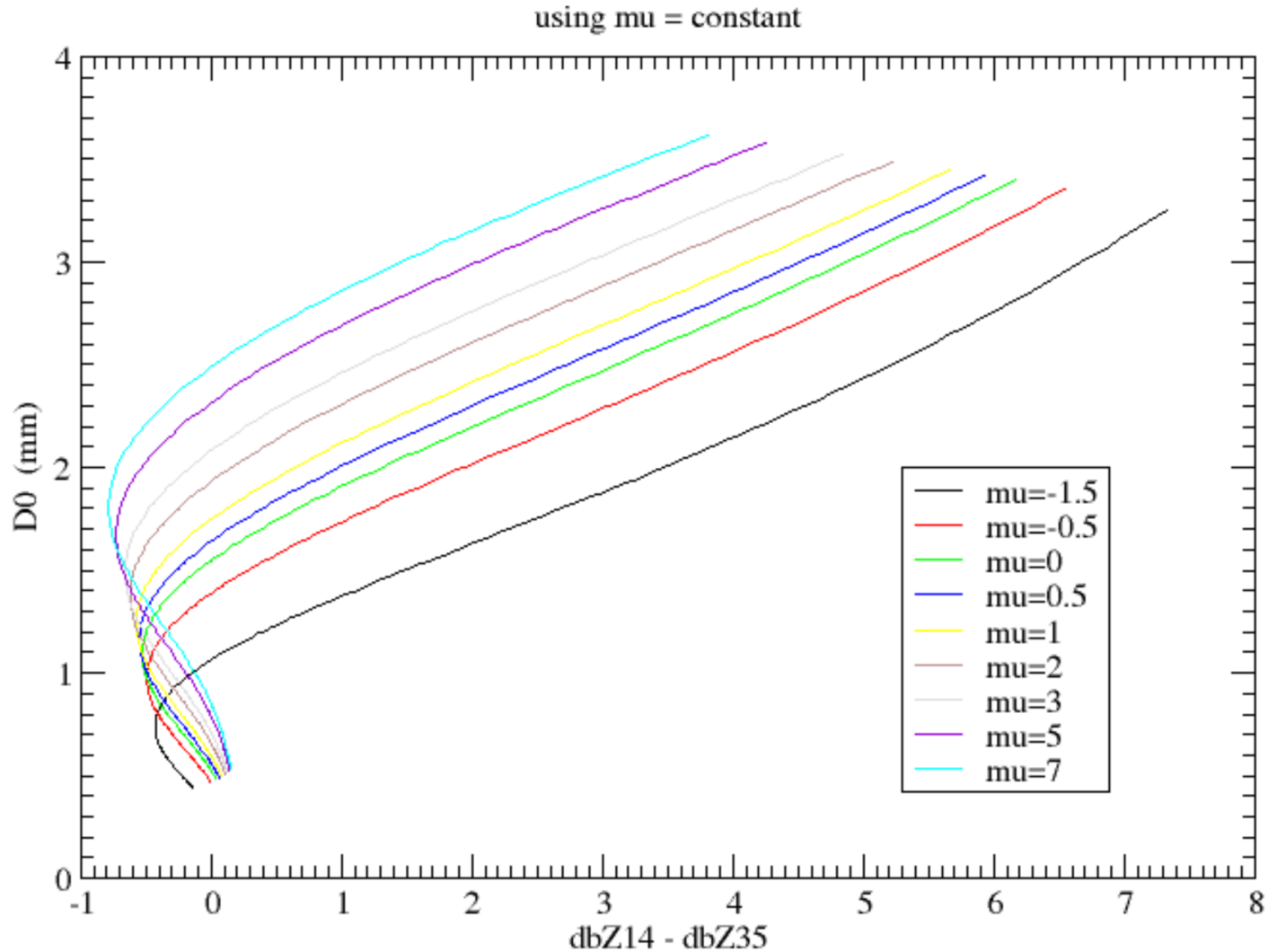


$$\sigma_m \neq \text{const } D_m^{0.5}$$

Darwin profiler, January 19+20 (blue), 22 (red) and 23 (green), 2006



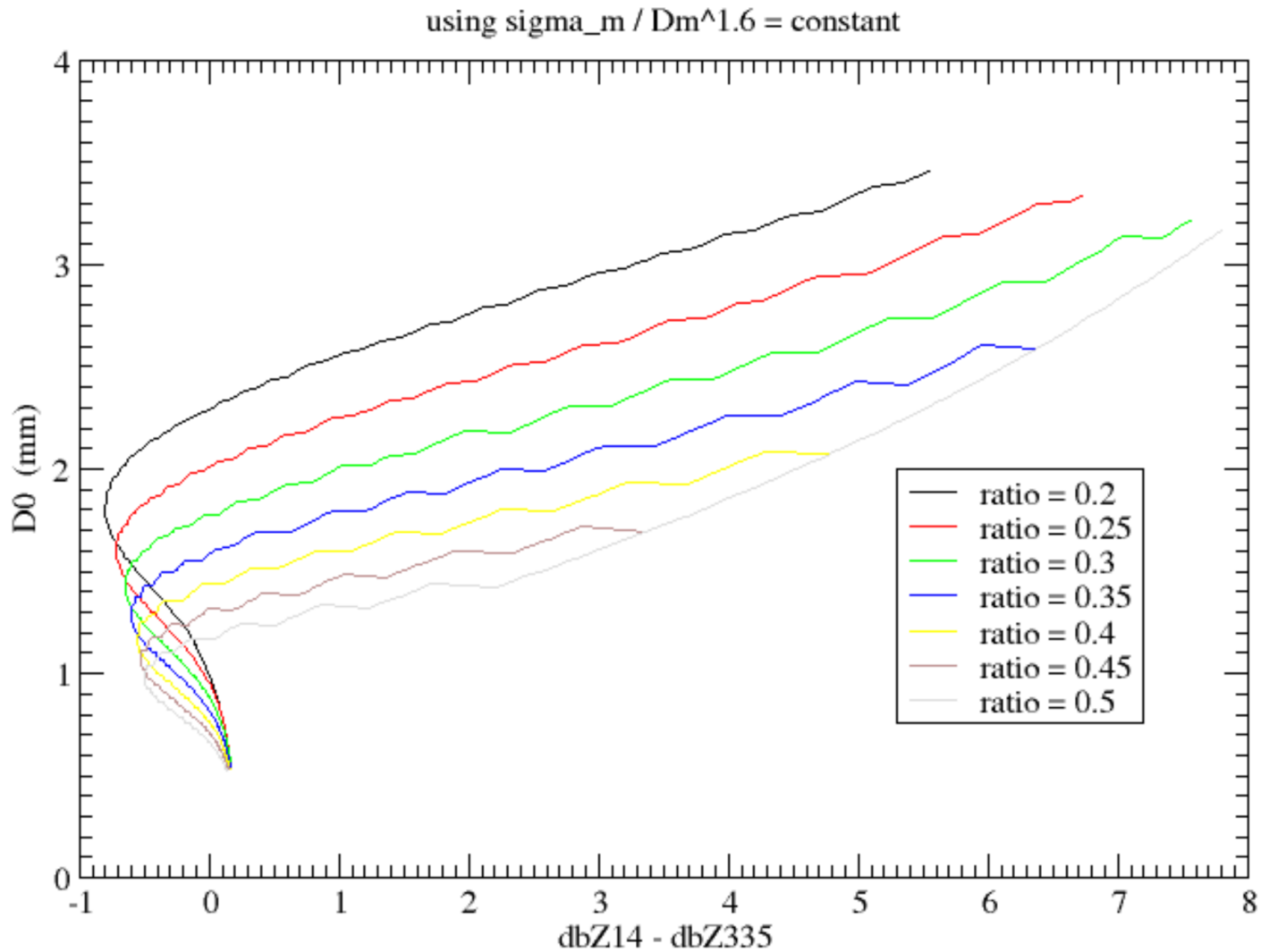
Coup de grâce: $\frac{\partial \mu}{\partial t} + V \cdot \nabla \mu = ?$ $\frac{\partial \sigma_m}{\partial t} + V \cdot \nabla \sigma_m = ?$

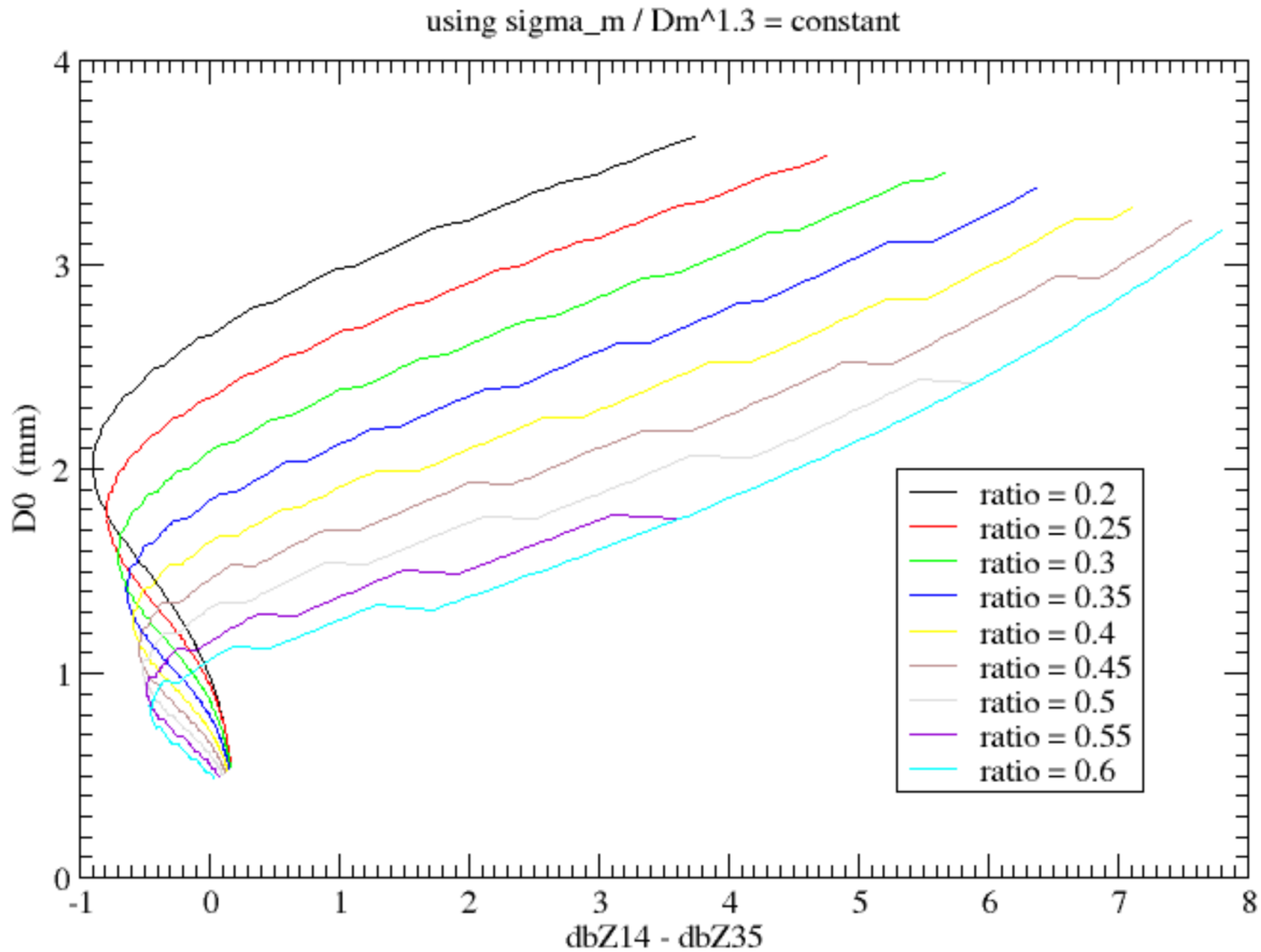




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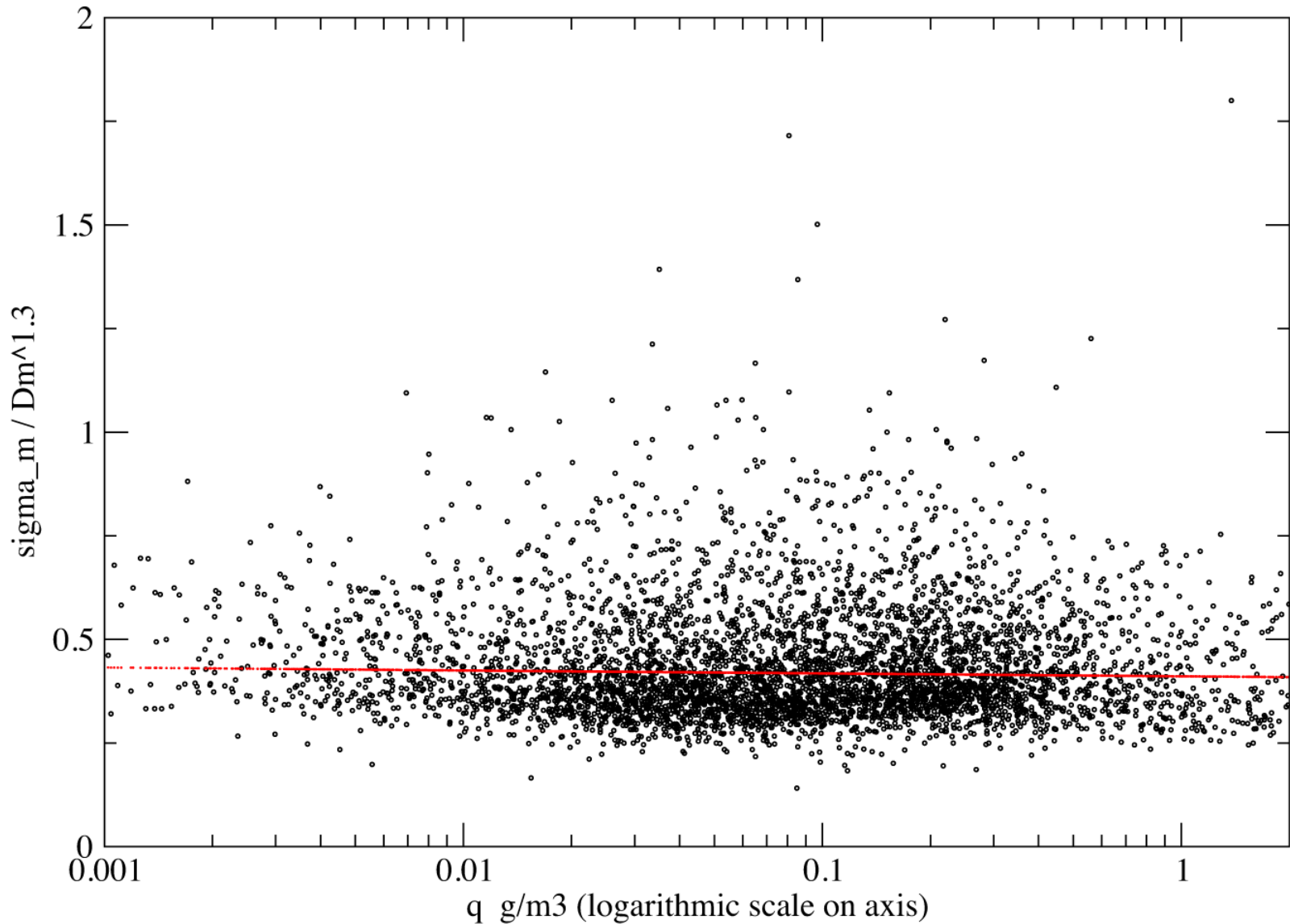


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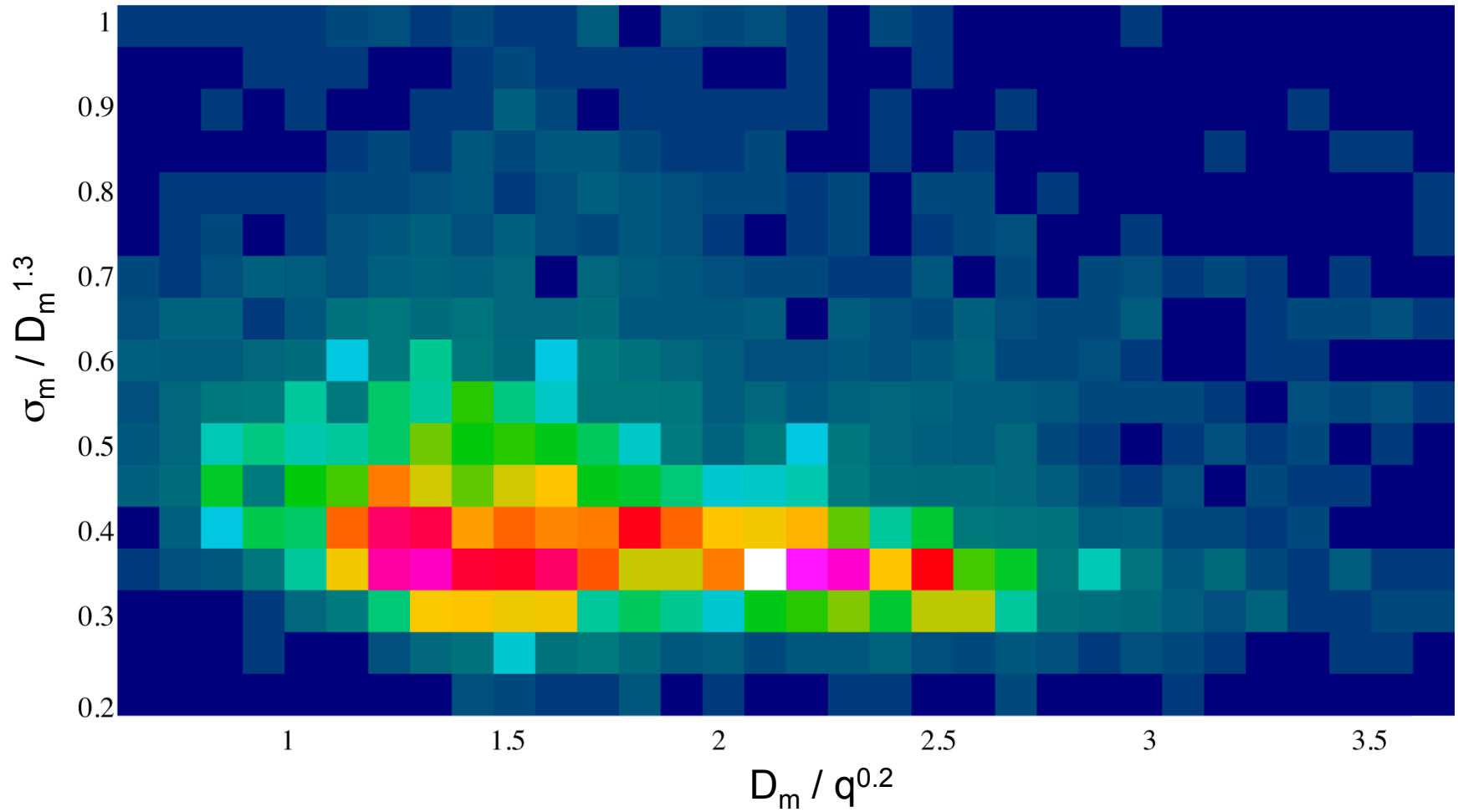
COARE





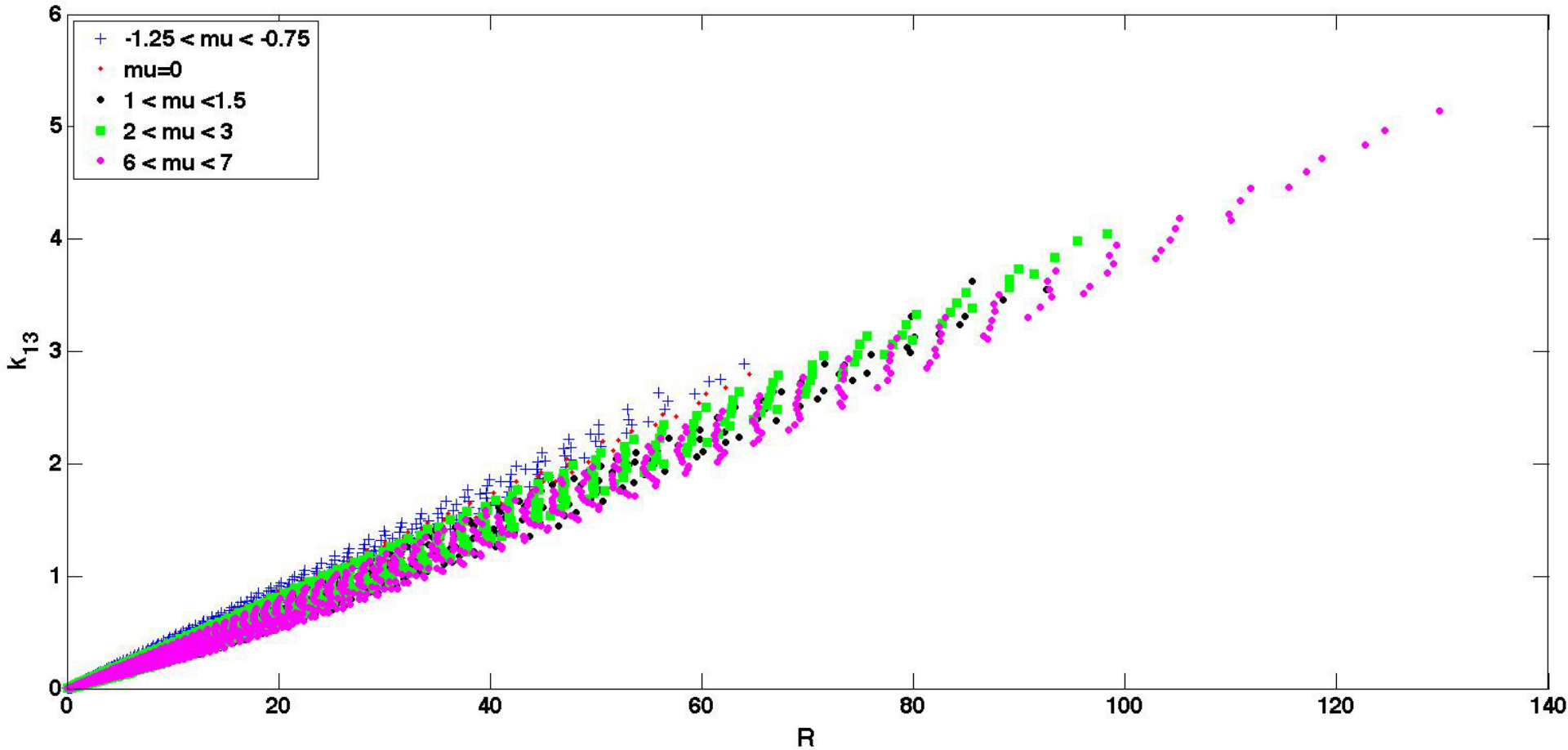
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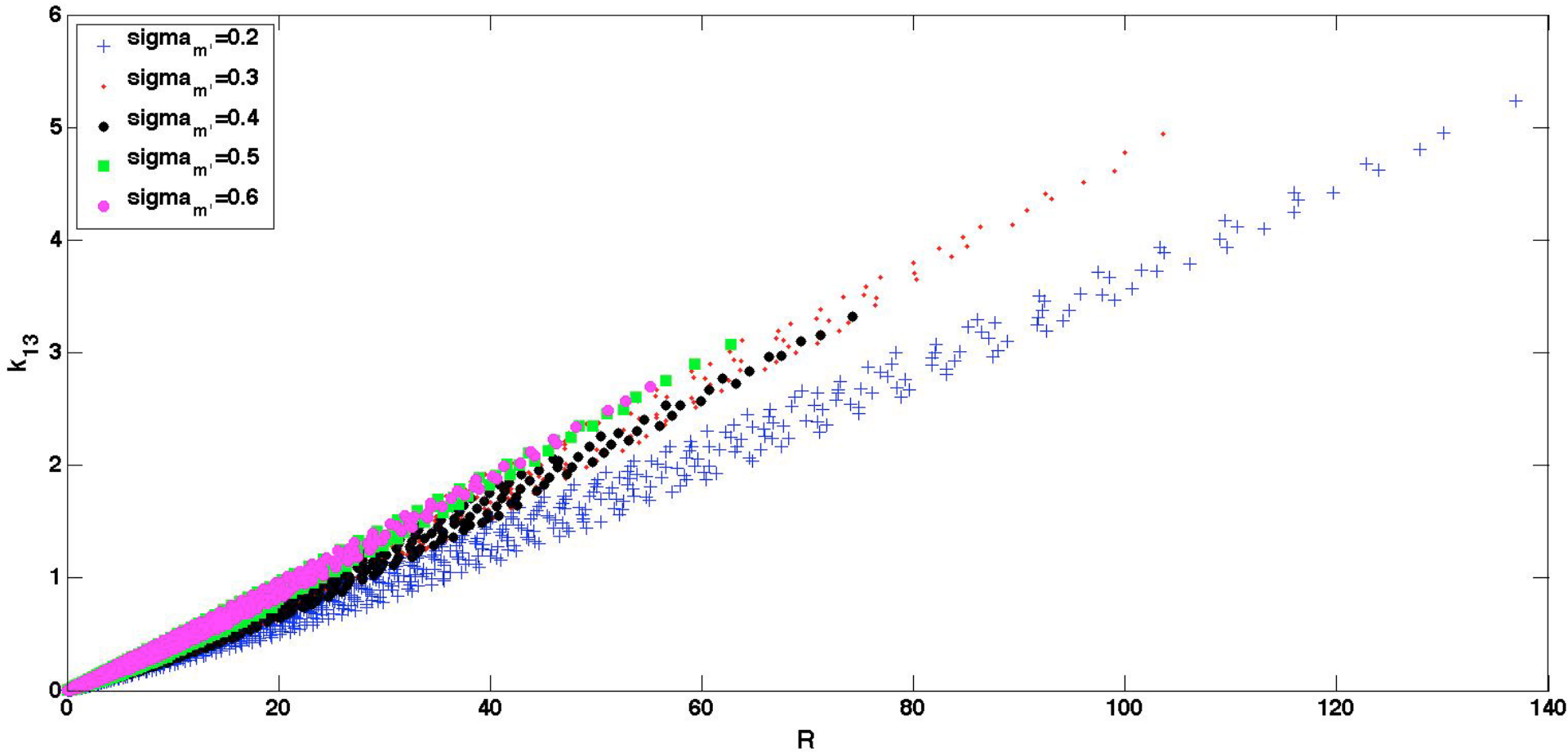


14GHz attenuation vs Rain rate -- μ constant assumption)



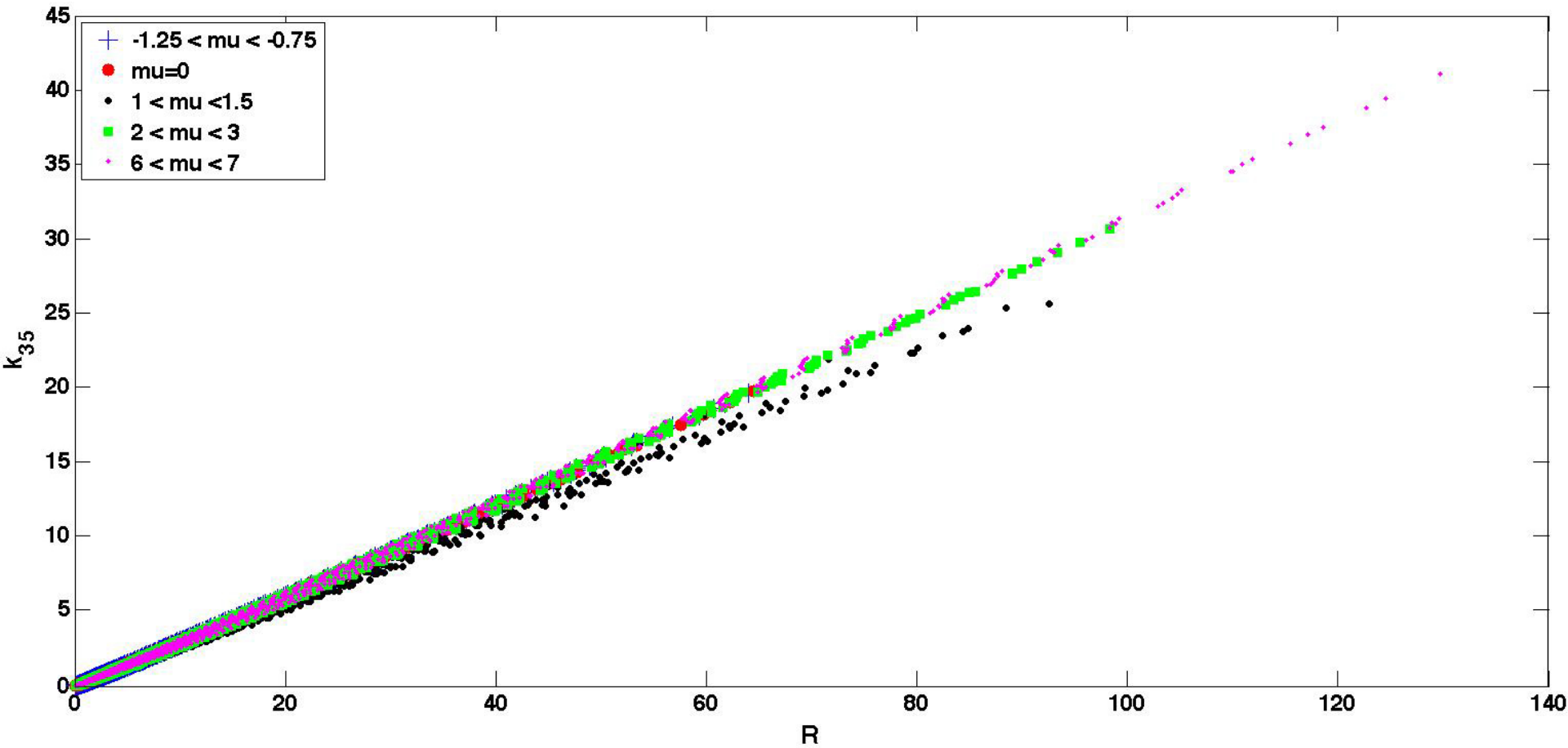


14GHz attenuation vs Rain rate -- $\sigma_m/D_m^{1.3}$ constant assumption)



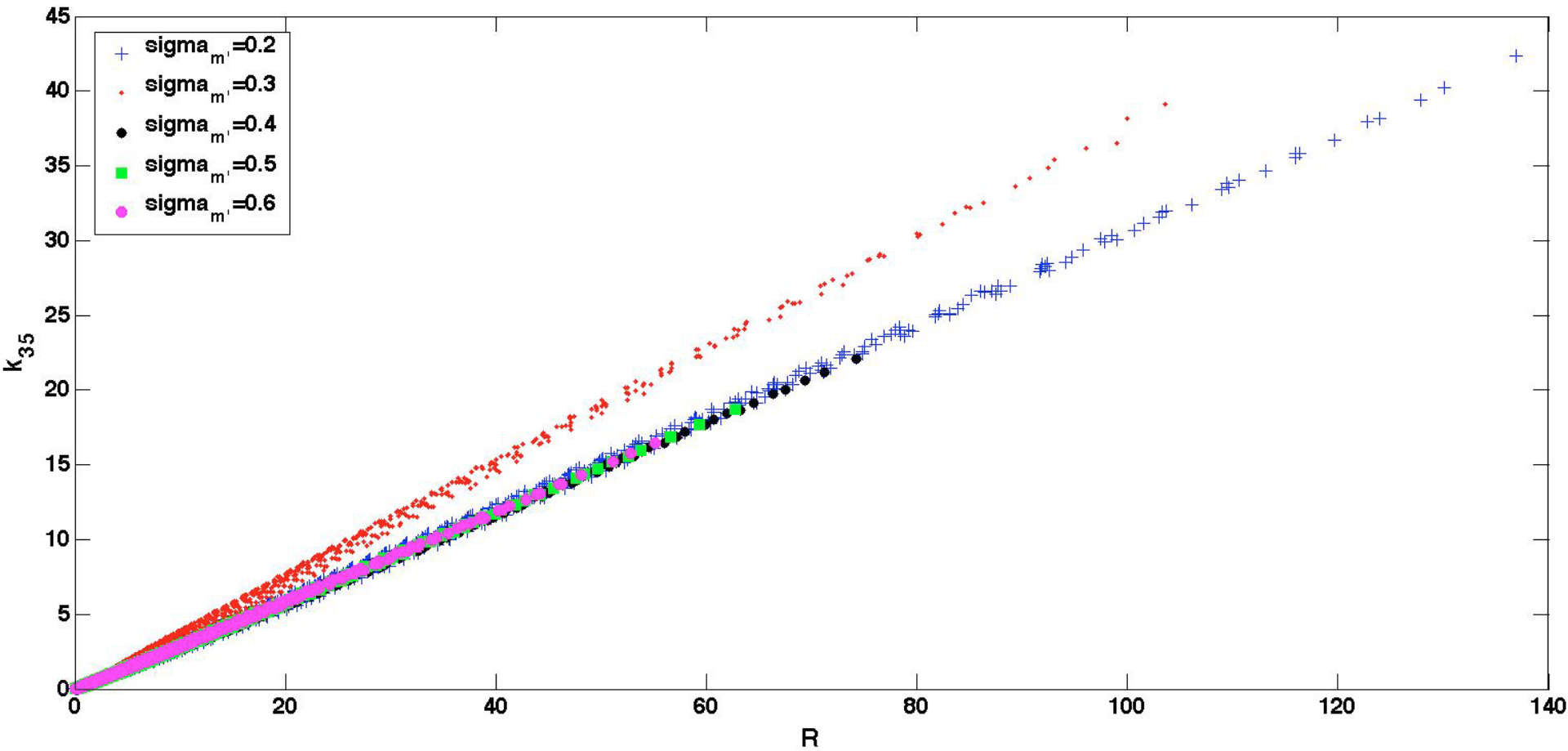


35GHz attenuation vs Rain rate -- μ constant assumption)



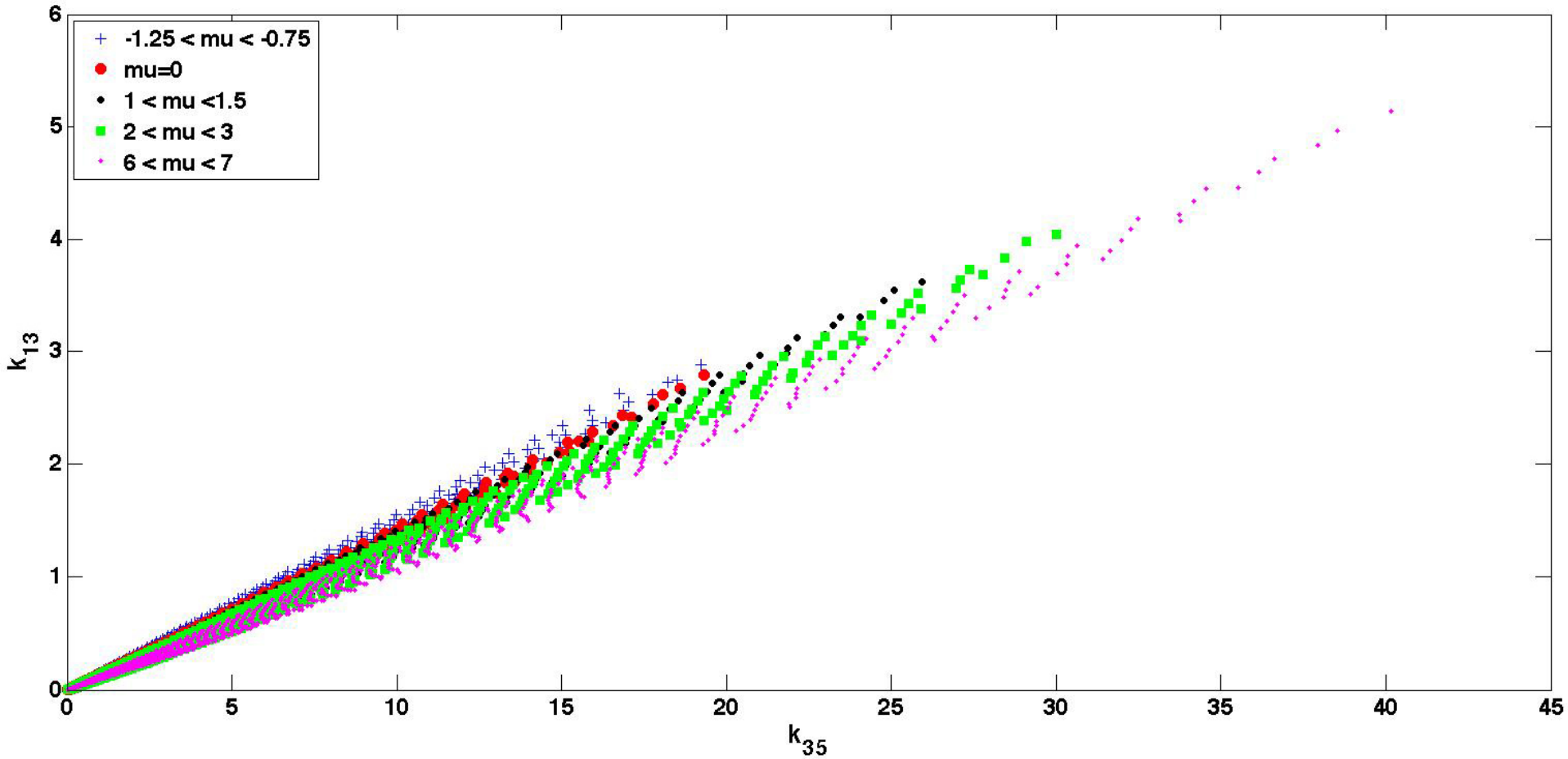


35GHz attenuation vs Rain rate -- $\sigma_m/D_m^{1.3}$ constant assumption)



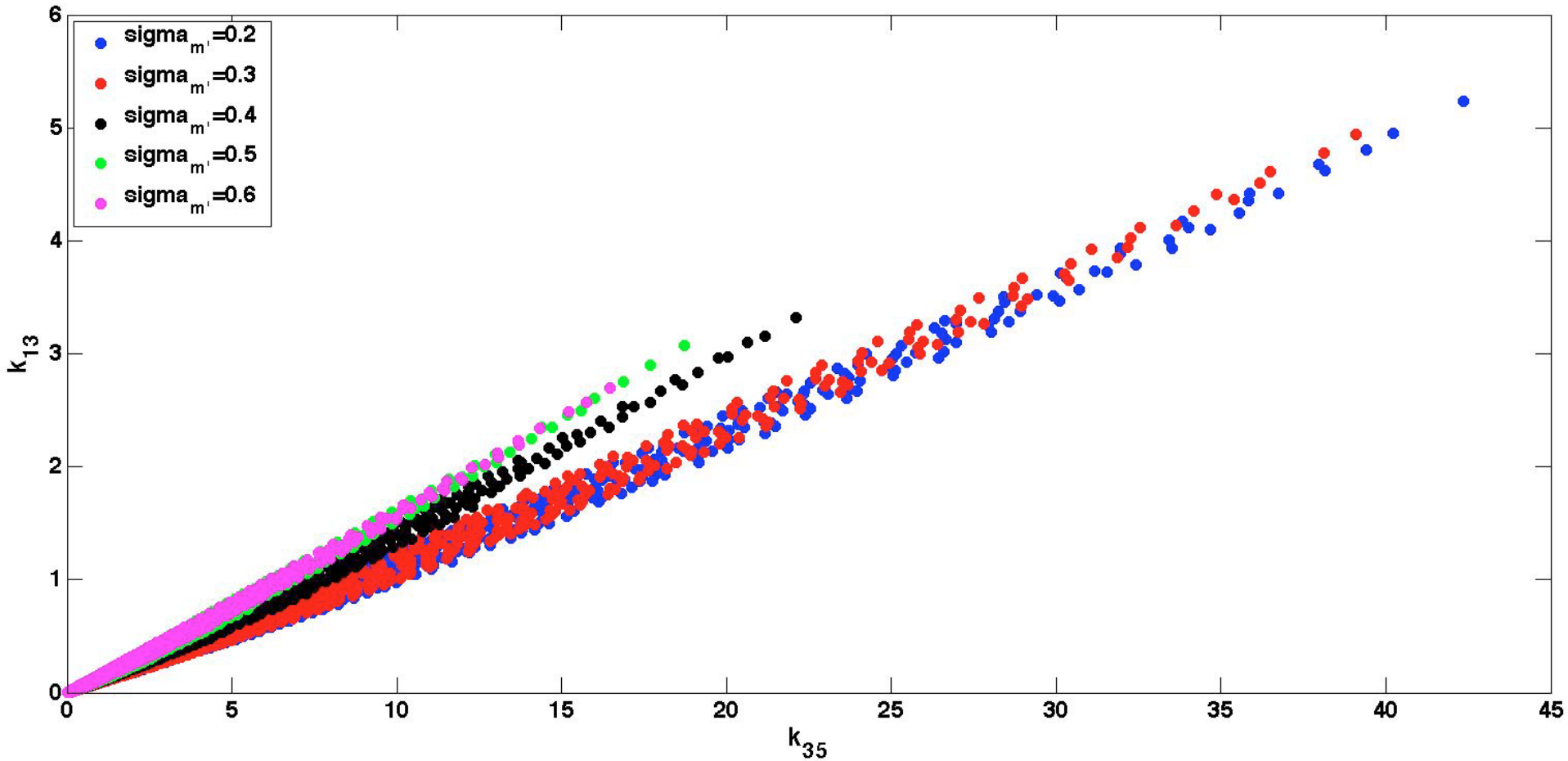


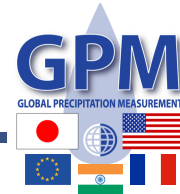
14GHz attenuation vs 35GHz attenuation -- μ constant assumption)





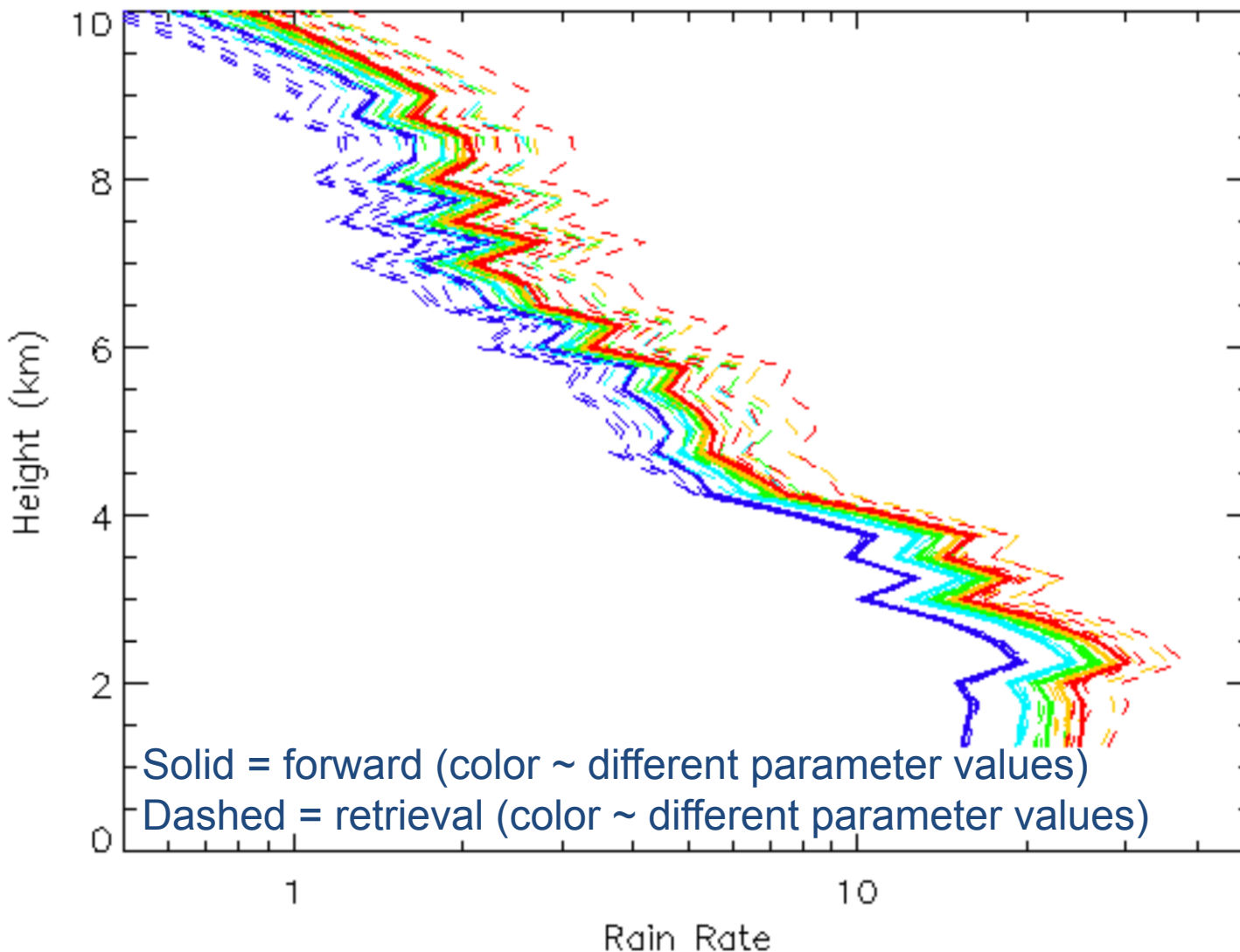
14GHz attenuation vs 35GHz attenuation -- $\sigma_m/D_m^{1.3}$ constant assumption)

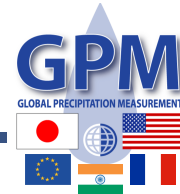




Algorithm testing: Start with TRMM, inject DSD (μ or $\sigma_m/D_m^{1.3}$) to generate Z35, then retrieve with DSD assumptions (μ or $\sigma_m/D_m^{1.3}$)

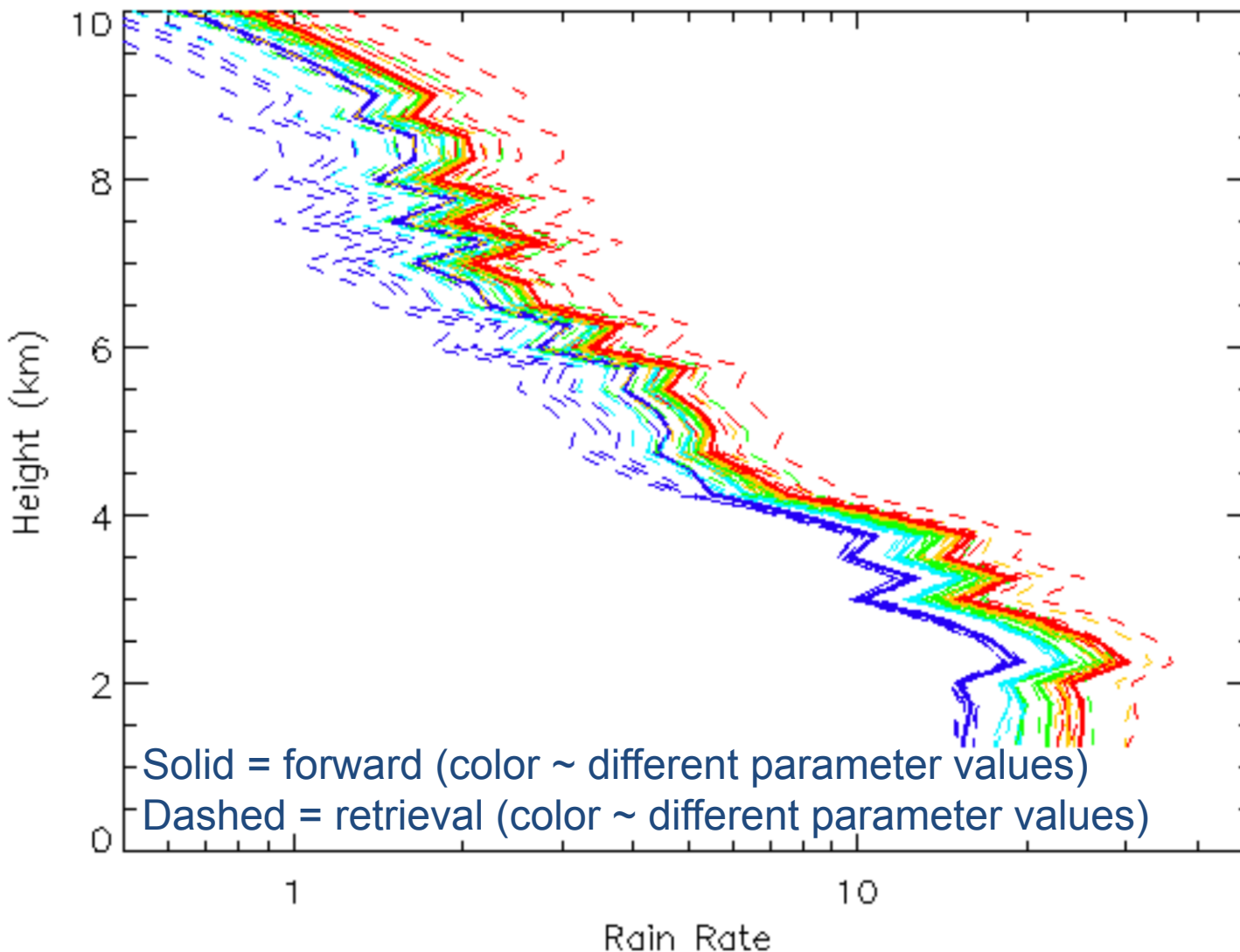
Mu Table with Mu Z35

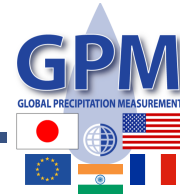




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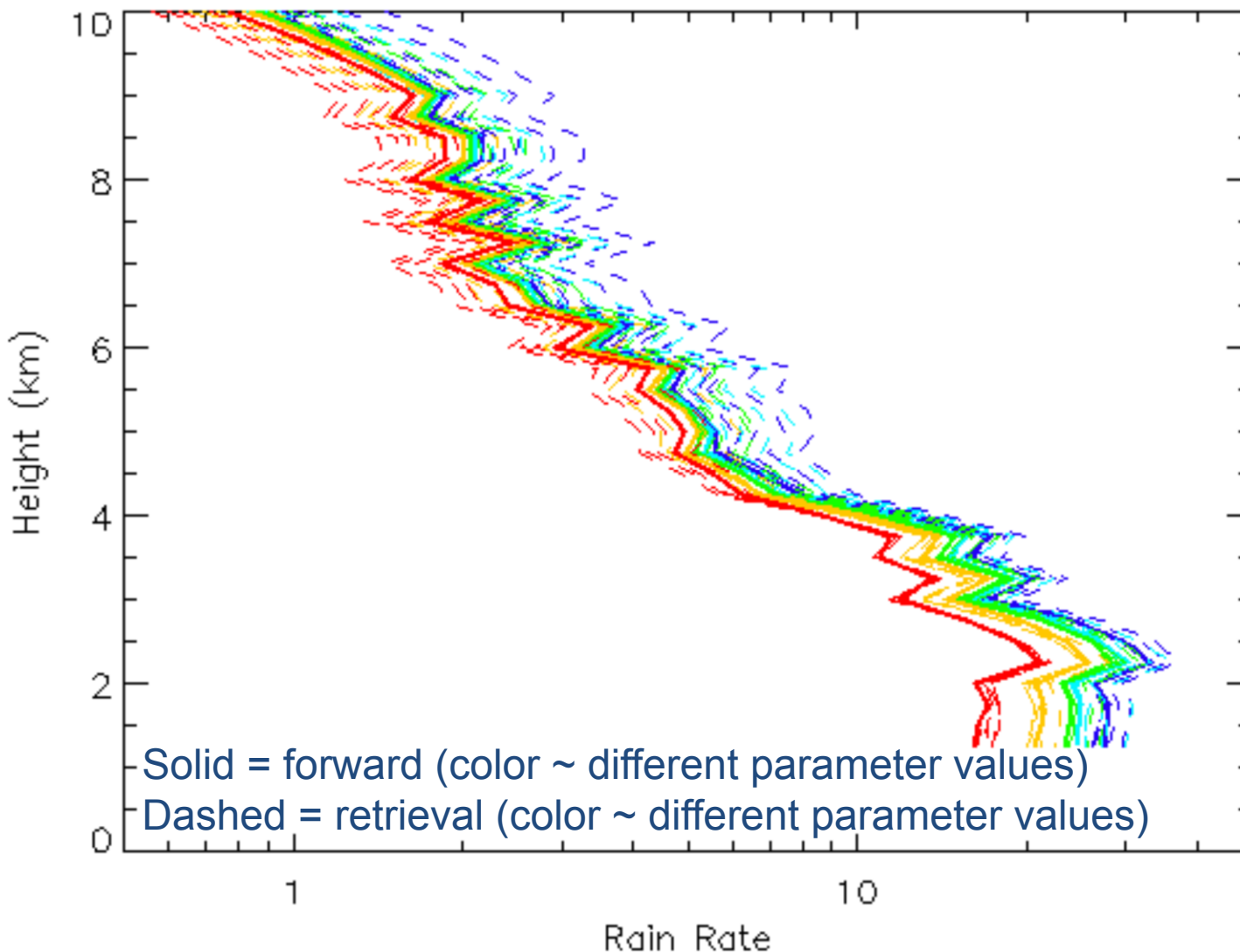
Ratio Table with Mu Z35





Algorithm testing: Start with TRMM, inject DSD (μ or $\sigma_m/D_m^{1.3}$) to generate Z35, then retrieve with DSD assumptions (μ or $\sigma_m/D_m^{1.3}$)

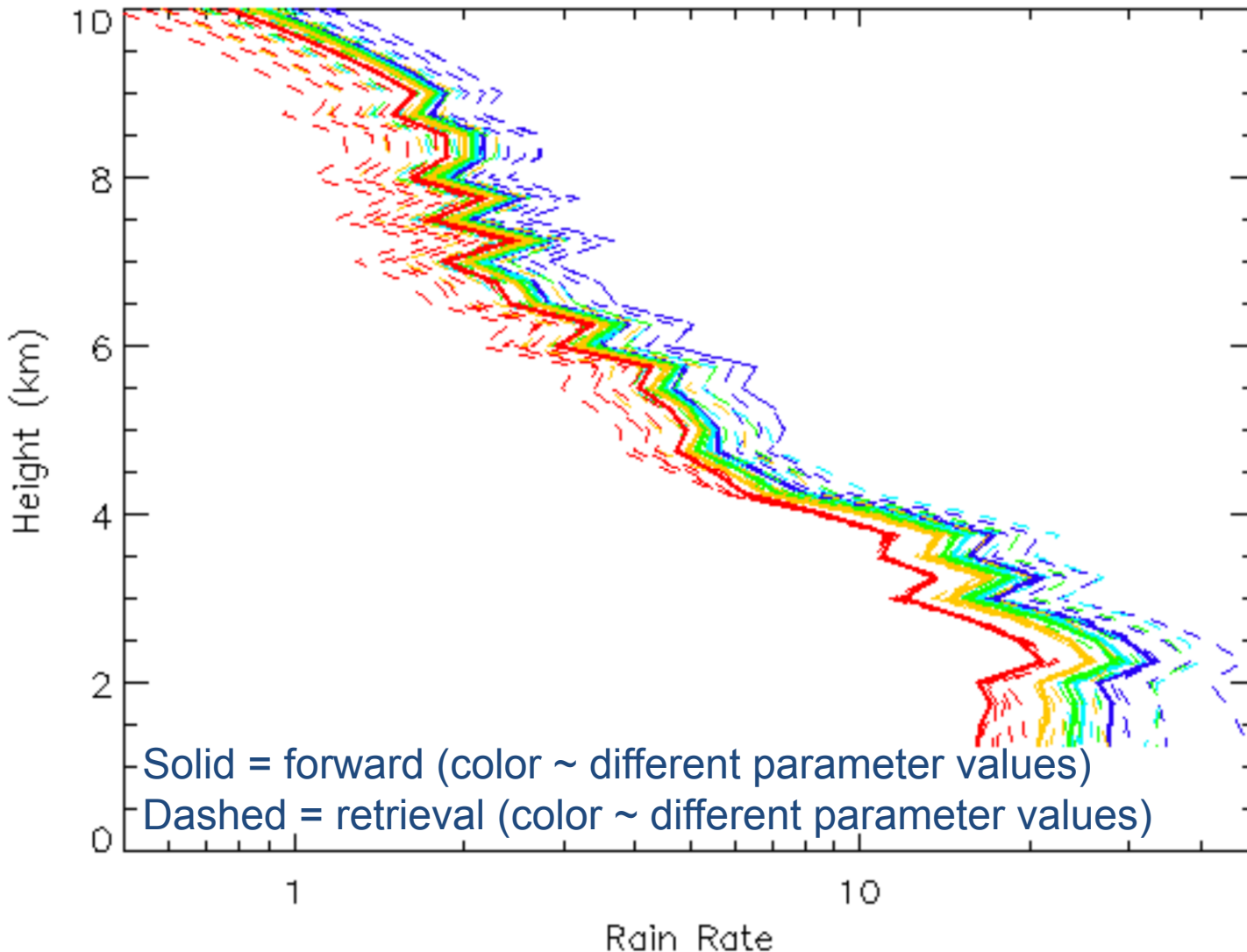
Mu Table with Ratio Z35





Algorithm testing: Start with TRMM, inject DSD (μ or $\sigma_m/D_m^{1.3}$) to generate Z35, then retrieve with DSD assumptions (μ or $\sigma_m/D_m^{1.3}$)

Ratio Table with Ratio Z35





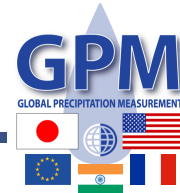
Smoothing Methodology (case of Tropical Cyclones):

- Start with WRF simulations (say HEDAS Earl 2010 h3vk, 2010-08-29-12Z to 2010-09-03-18Z), using stream ψ , potential χ , P , T , RH , W , q_{cliq} , q_r , q_{cli} , q_s , q_g , q_h at 42 vertical levels for a total of 504 variables x_1, \dots, x_{504}
- for each of these 12million columns, forward-calculate T_{b1}, \dots, T_{b9}
- find the principal components x_1', \dots, x_{504}' (each is a linear combo of x_1, \dots, x_{504}) and the principal components T_1', \dots, T_9' (each a linear combo of T_1', \dots, T_9')
- Then we will have to find combos of x_1', \dots, x_{504}' that correlate most with combos of T_1', \dots, T_9'
- Say these combos are x_1'', x_2'', x_3'' and T_1'', T_2'', T_3'' : we finally need to express the latter in terms of the former, in a differentiable way (to be able to compute derivatives)



Smoothing Methodology (case of Tropical Cyclones):

- Start with WRF simulations (say HEDAS Earl 2010 h3vk, 2010-08-29-12Z to 2010-09-03-18Z), using stream ψ , potential χ , P , T , RH , W , q_{cliq} , q_r , q_{cli} , q_s , q_g , q_h at 42 vertical levels for a total of 504 variables x_1, \dots, x_{504}
- for each of these 12million columns, forward-calculate T_{b1}, \dots, T_{b9}
- **Step 1:** find the principal components x_1', \dots, x_{504}'
- **Step 2:** find the principal components T_1', \dots, T_9'
- **Step 3:** find combos of x_1', \dots, x_{504}' that correlate most with combos of T_1', \dots, T_9'



Smoothing Methodology (case of Tropical Cyclones):

- Start with HWRP simulations (say HEDAS Earl 2010 h3vk, 2010-08-29-12Z to 2010-09-03-18Z), using stream ψ , potential χ , P, T, RH, W, q_{cliq} , q_r , q_{cli} , q_s , q_g , q_h at 42 vertical levels for a total of 504 variables x_1, \dots, x_{504}
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• **Step 2:** find the principal components T_1', \dots, T_9'

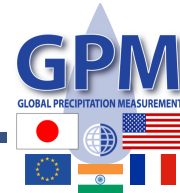
• **Step 3:** find

combos of x_1', \dots, x_{504}' that correlate most with combos of T_1', \dots, T_9'

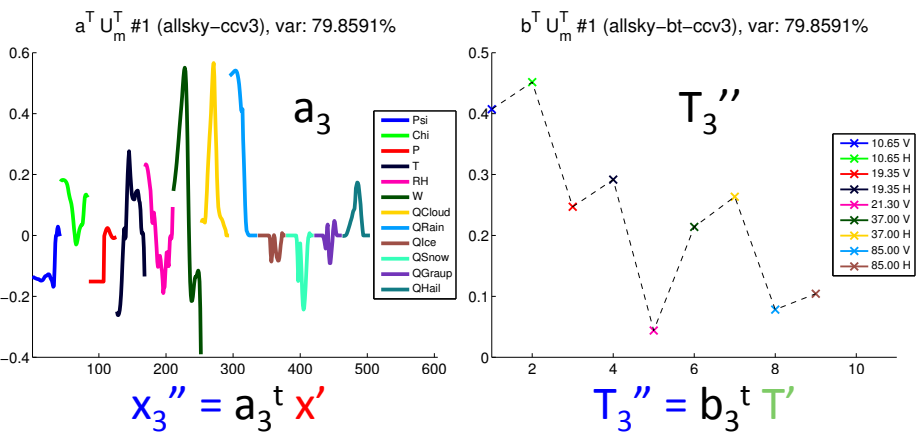
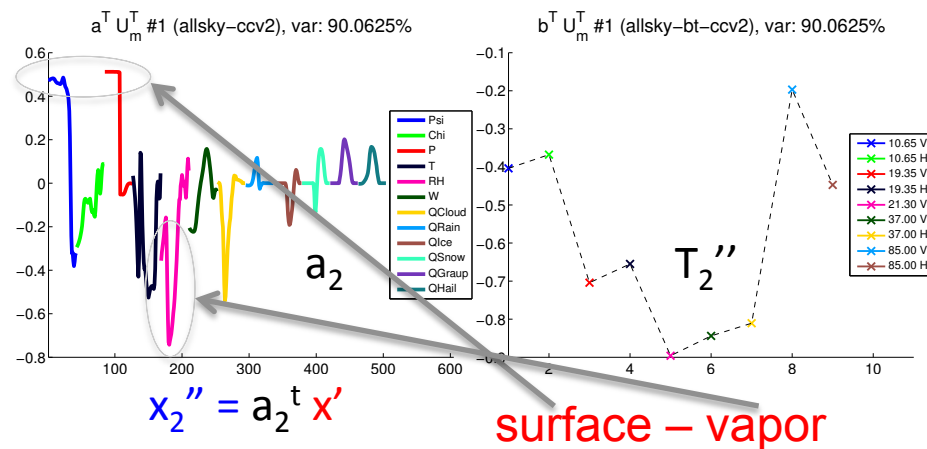
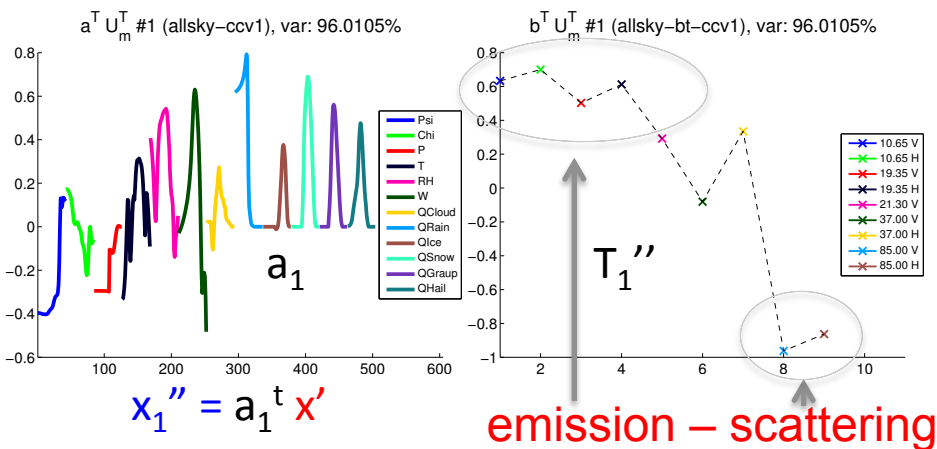
and express T_1'', T_2'', T_3'' differentiably in terms of x_1'', x_2'', x_3''

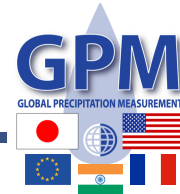
$$T_i''(x_1'', x_2'', x_3'') = \sum T_i''^{(n)} \exp(-[x_1''-x_1''^{(n)}]^2 -[x_2''-x_2''^{(n)}]^2 -[x_3''-x_3''^{(n)}]^2)$$

where the weighted sum over n runs over the 12million training points

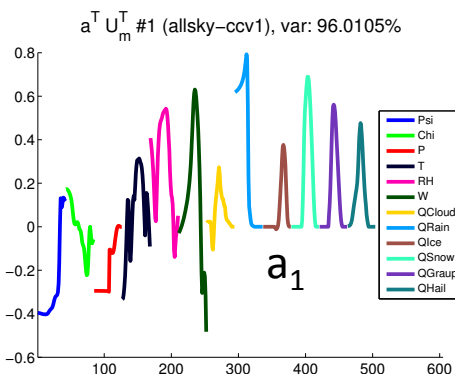


First part of step 3: here are the first 3 x'' and T''

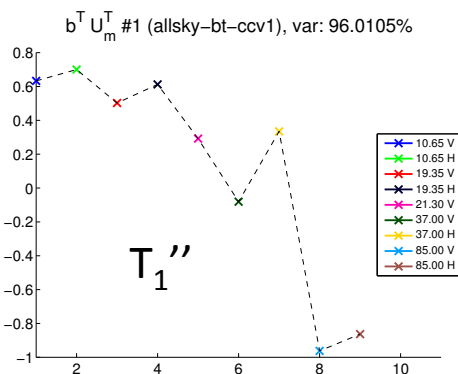




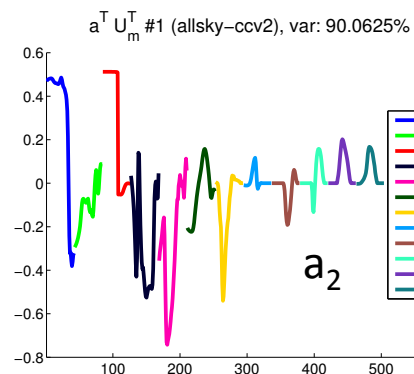
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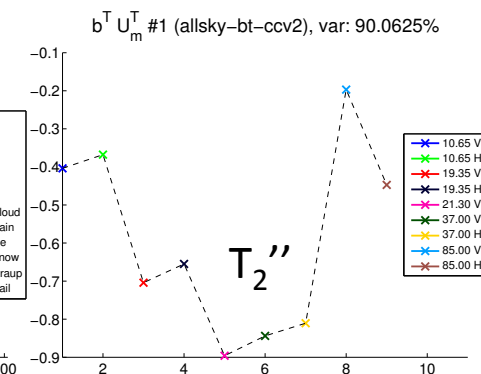
$$x_1'' = a_1^t x'$$



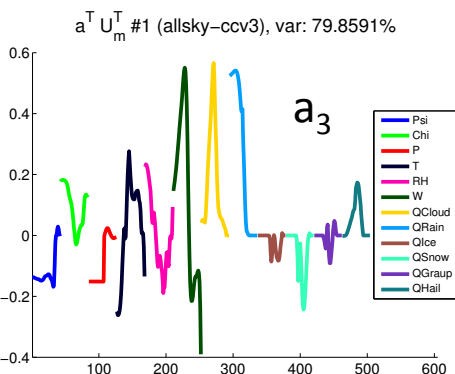
emission – scattering



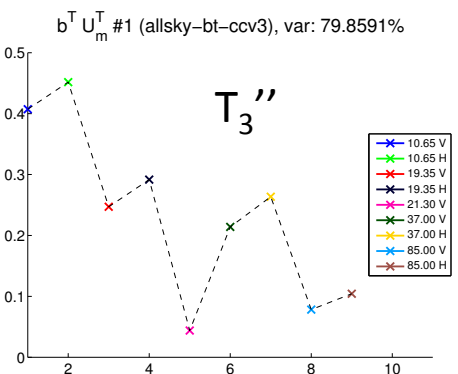
$$x_2'' = a_2^t x'$$



surface – vapor



$$x_3'' = a_3^t x'$$



$$T_3'' = b_3^t T'$$

Most remarkable:
the operators H_1, H_2, H_3
giving

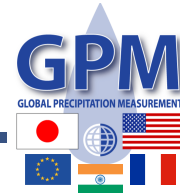
$$T_1'' = H_1(x_1'', x_2'', x_3'')$$

$$T_2'' = H_2(x_1'', x_2'', x_3'')$$

$$T_3'' = H_3(x_1'', x_2'', x_3'')$$

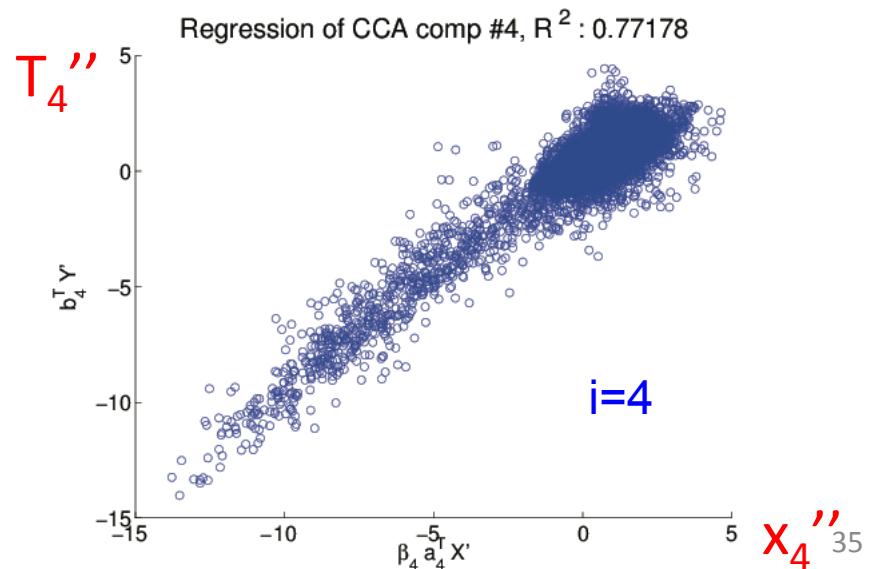
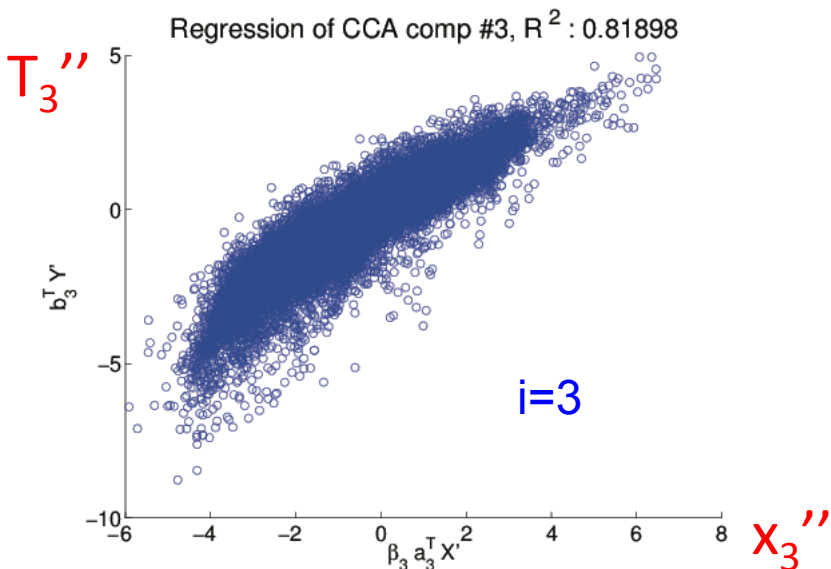
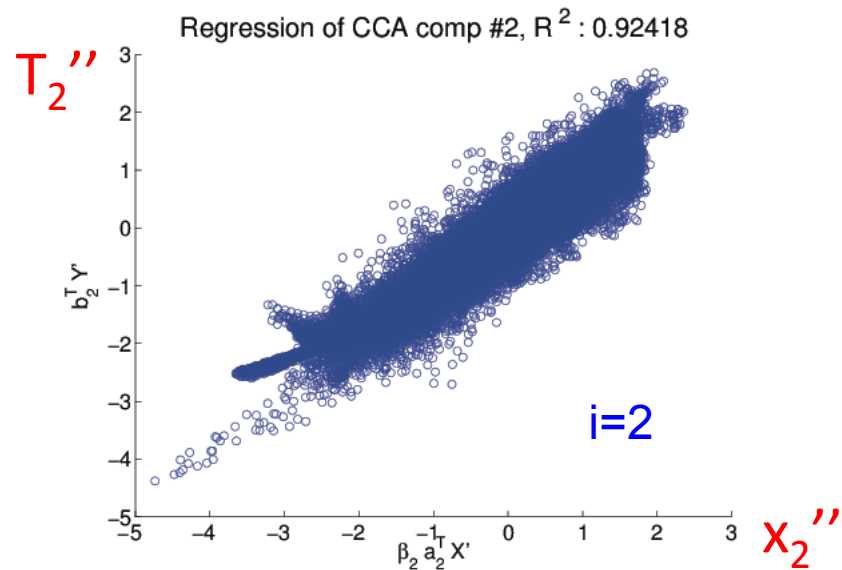
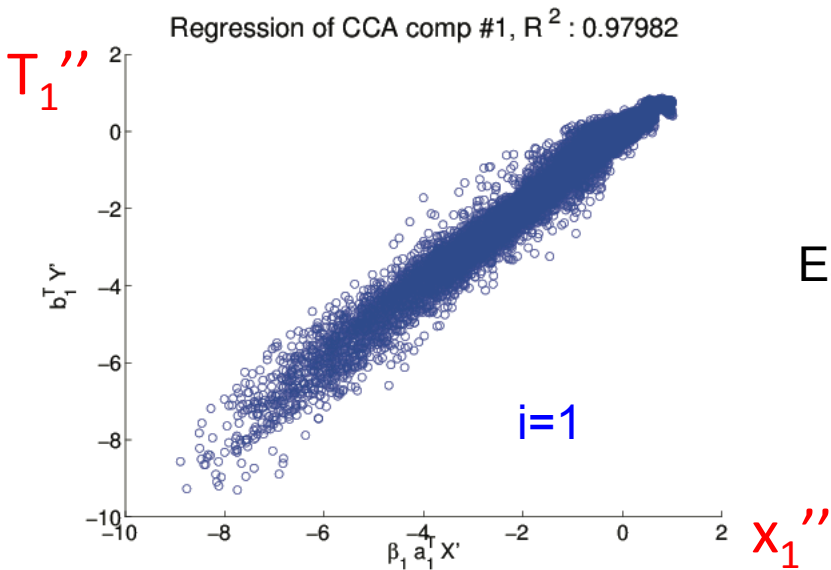
are not so nonlinear:





First part of step 3: T_i'' (vertical) vs x_i'' (horizontal)

Earl 2010





Second part of step 3: use nonlinear expression

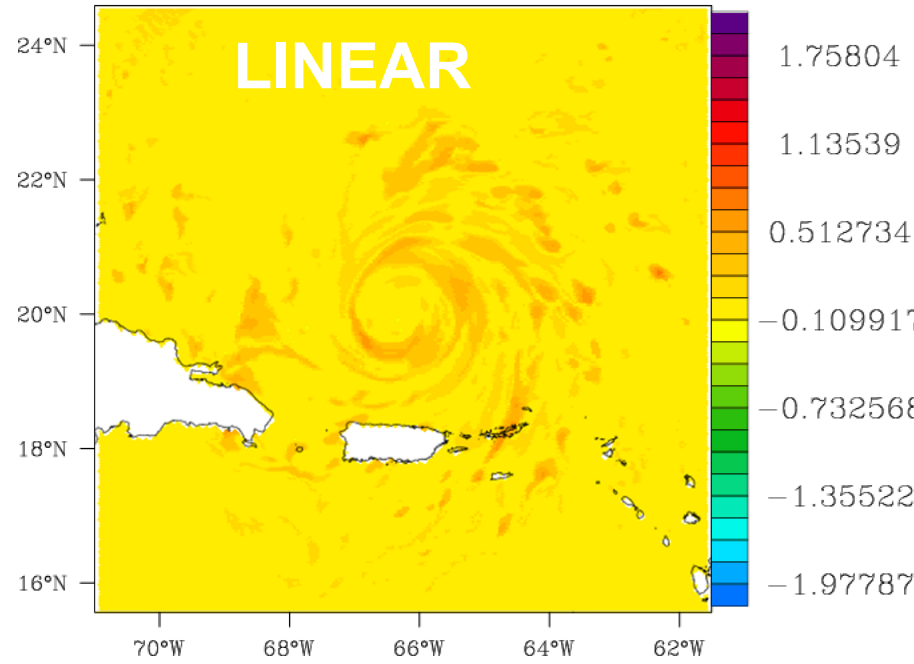
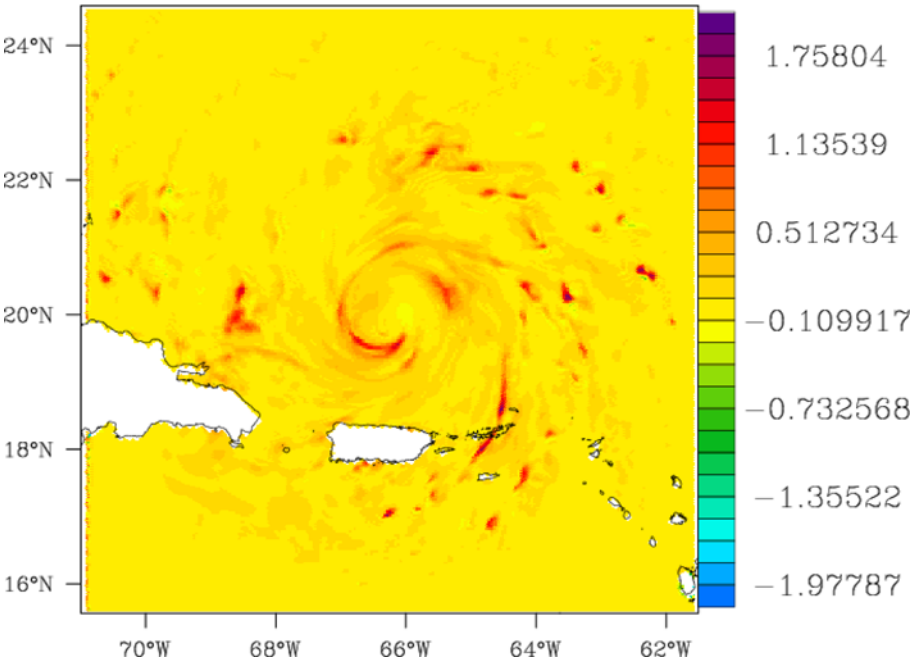
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So let's try an assimilation using this observation operator:
(start with 1D-var – not quite a “retrieval”, because of the covariances)

Having started with a horizontally uniform background, each variable having the global mean value at that height level:

Avg w levels 0-41, truth (m/s)

Avg w levels 0-41, anlys (m/s)



vertical component of wind



Second part of step 3: use nonlinear expression

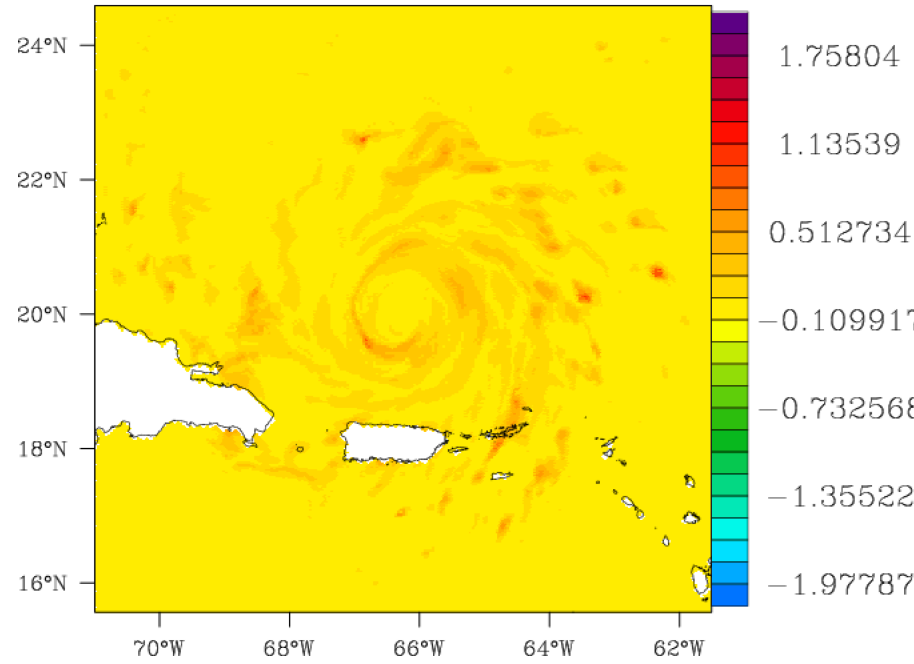
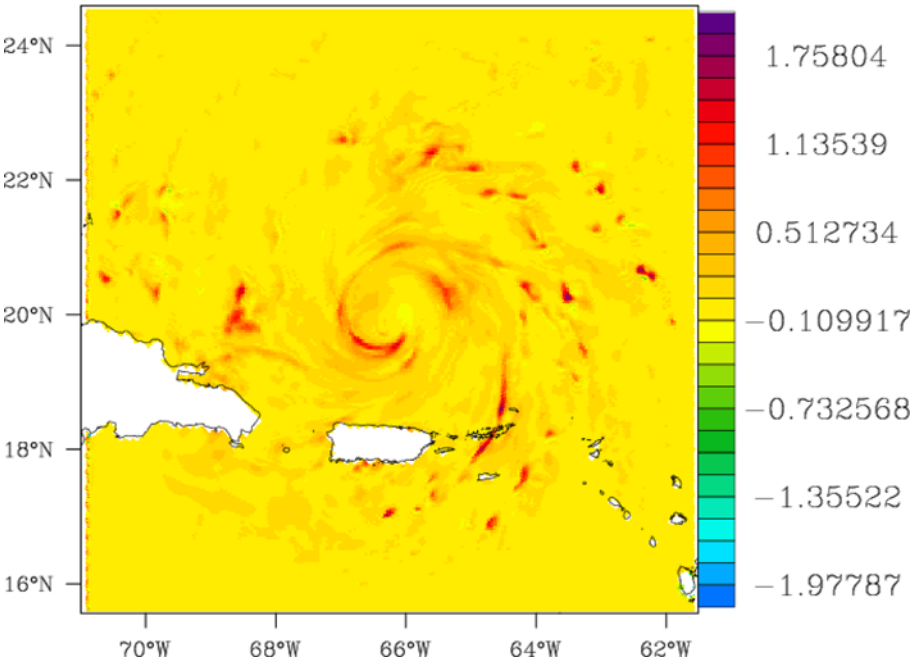
$$T_i''(x_1'', x_2'', x_3'') = \sum T_i''^{(n)} \exp(-[x_1''-x_1''^{(n)}]^2 -[x_2''-x_2''^{(n)}]^2 -[x_3''-x_3''^{(n)}]^2)$$

So let's try an assimilation using this observation operator:
(start with 1D-var – not quite a “retrieval”, because of the covariances)

Having started with a horizontally uniform background, each variable having the global mean value at that height level:

Avg w levels 0-41, truth (m/s)

Avg w levels 0-41, anlys (m/s)



vertical component of wind



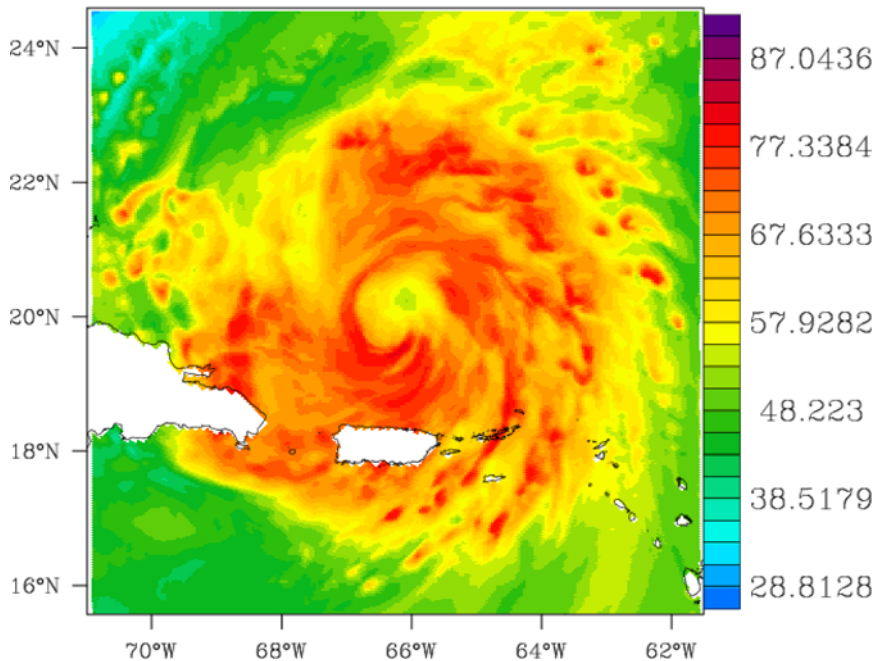
Second part of step 3: use nonlinear expression

$$T_i'' \sim x_i''$$

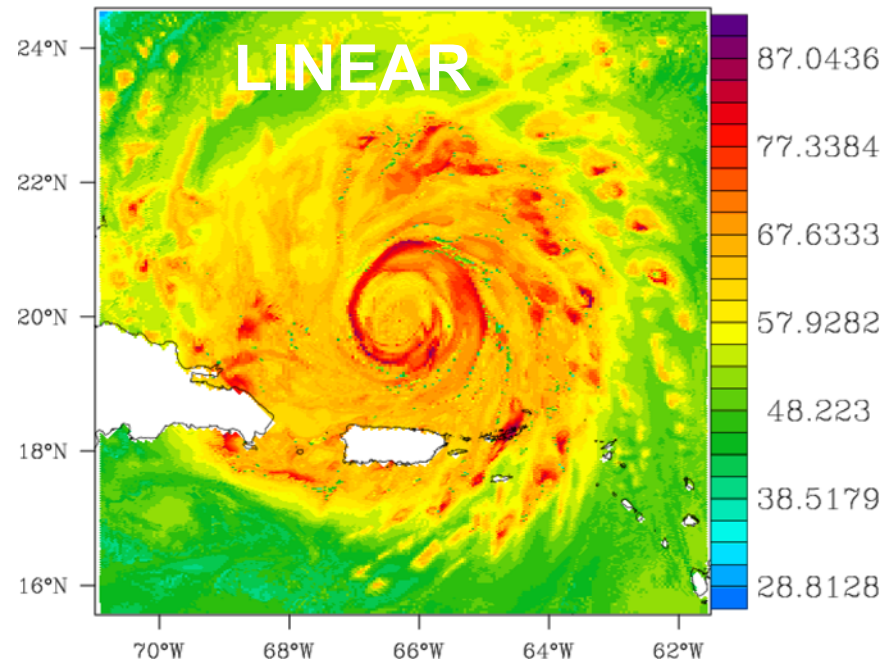
So let's try an assimilation using this observation operator:
(start with 1D-var – not quite a “retrieval”, because of the covariances)

Having started with a horizontally uniform background, each variable having the global mean value at that height level:

Avg rh levels 0-41, truth (%)



Avg rh levels 0-41, anlys (%)



water vapor

Second part of step 3: use nonlinear expression

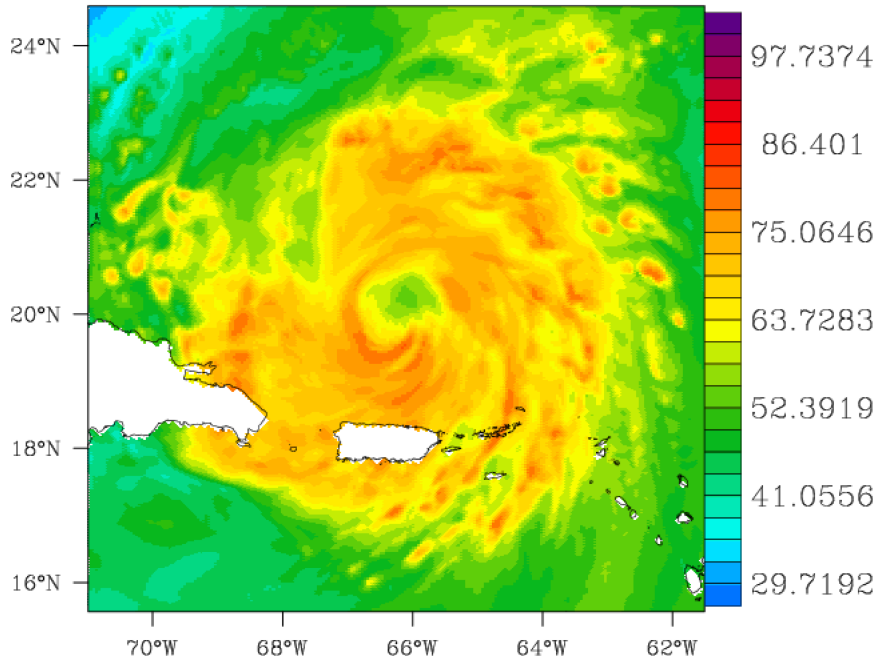
$$T_i''(x_1'', x_2'', x_3'') = \sum T_i''^{(n)} \exp(-[x_1'' - x_1''^{(n)}]^2 - [x_2'' - x_2''^{(n)}]^2 - [x_3'' - x_3''^{(n)}]^2)$$

So let's try an assimilation using this observation operator:

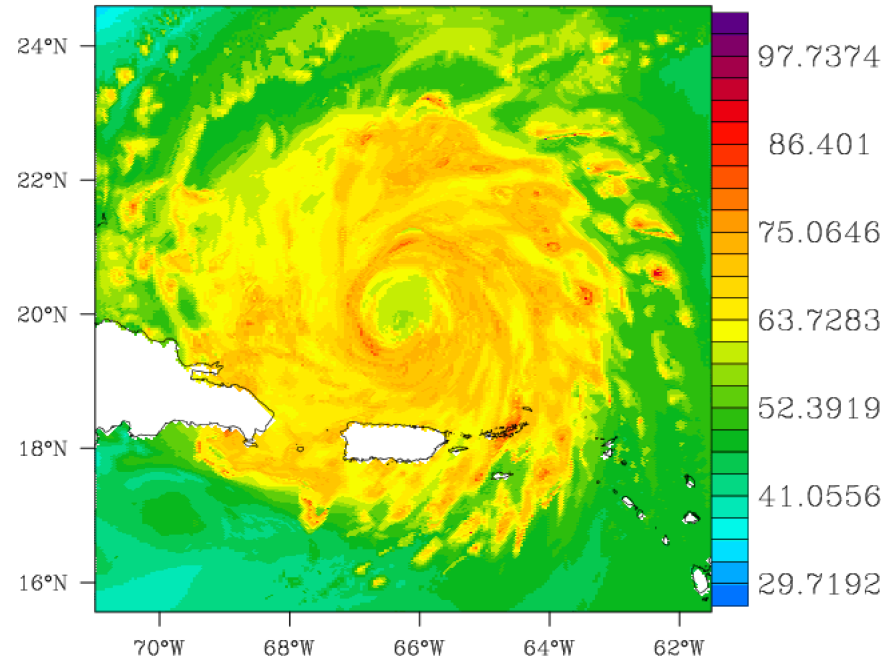
(start with 1D-var – not quite a “retrieval”, because of the covariances)

Having started with a horizontally uniform background, each variable having the global mean value at that height level:

Avg rh levels 0-41, truth (%)



Avg rh levels 0-41, anlys (%)





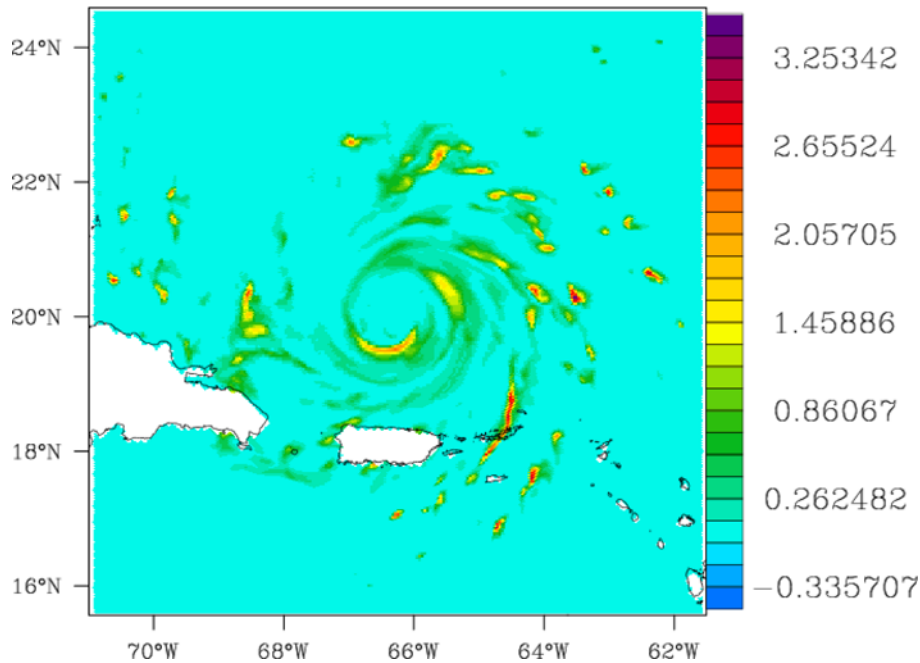
Second part of step 3: use nonlinear expression

$$T_i'' \sim x_i''$$

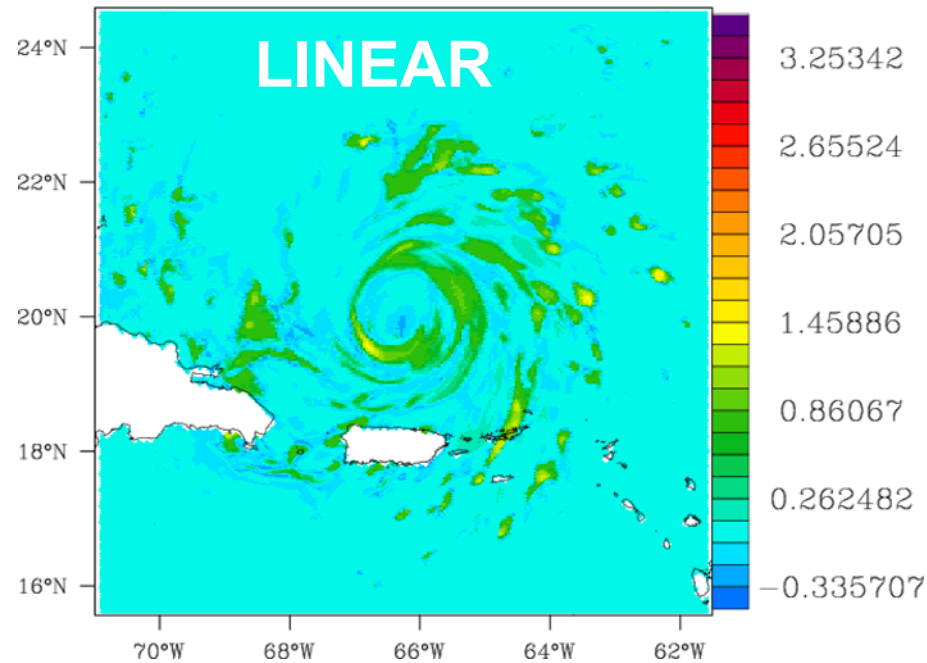
So let's try an assimilation using this observation operator:
(start with 1D-var – not quite a “retrieval”, because of the covariances)

Having started with a horizontally uniform background, each variable having the global mean value at that height level:

Avg grain levels 0-41, truth (g/kg)



Avg grain levels 0-41, anlys (g/kg)



RAIN



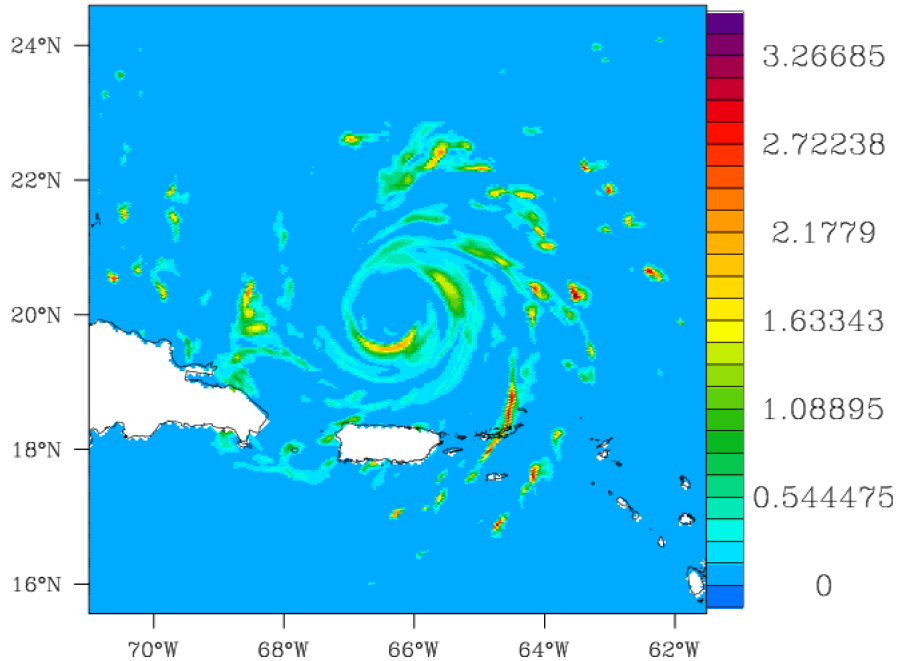
Second part of step 3: use nonlinear expression

$$T_i''(x_1'', x_2'', x_3'') = \sum T_i''^{(n)} \exp(-[x_1''-x_1''^{(n)}]^2 -[x_2''-x_2''^{(n)}]^2 -[x_3''-x_3''^{(n)}]^2)$$

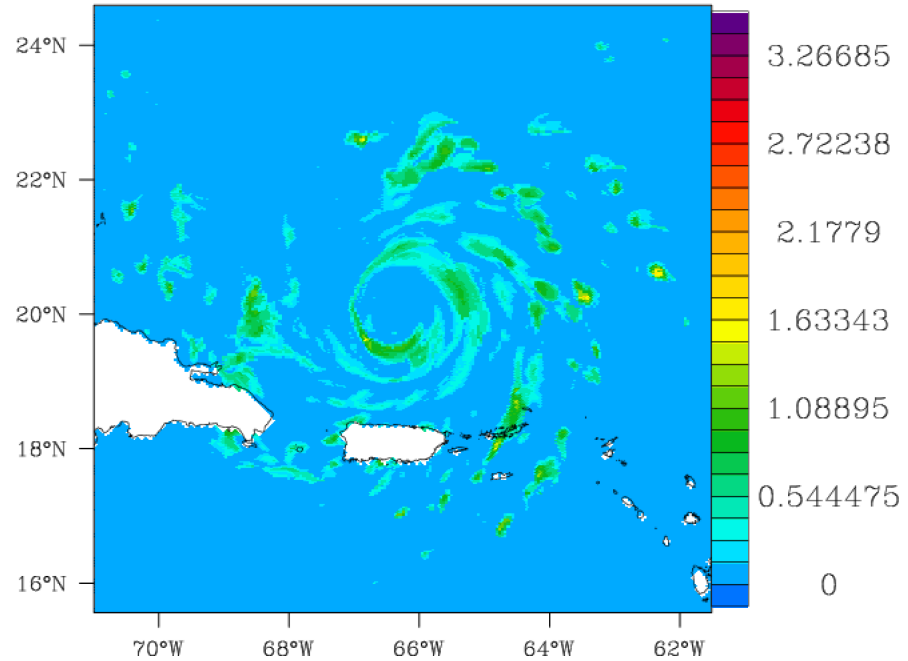
So let's try an assimilation using this observation operator:
(start with 1D-var – not quite a “retrieval”, because of the covariances)

Having started with a horizontally uniform background, each variable having the global mean value at that height level:

Avg grain levels 0-41, truth (g/kg)



Avg grain levels 0-41, anlys (g/kg)



RAIN

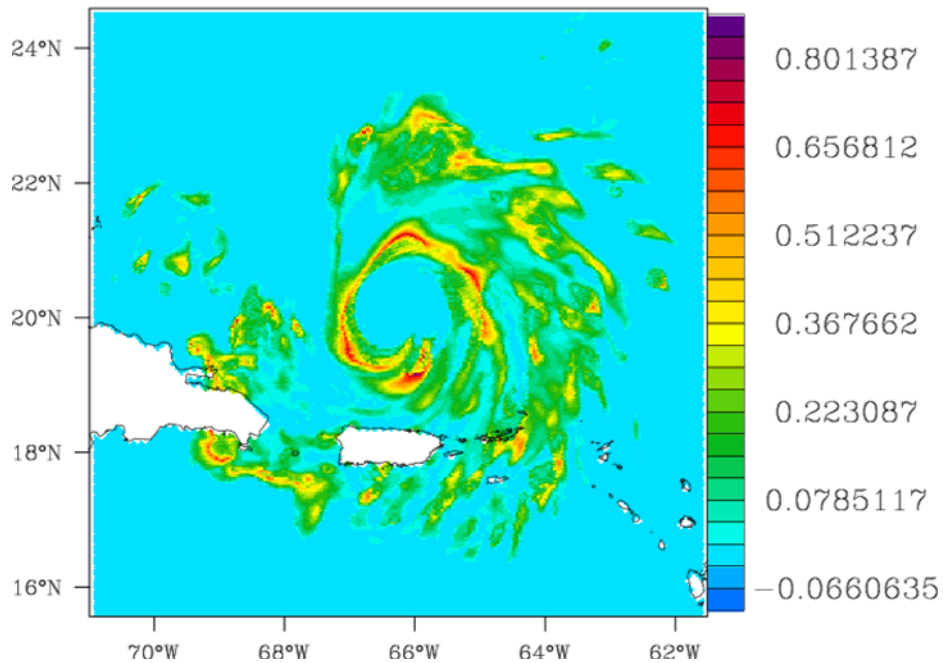
Second part of step 3: use nonlinear expression

$$T_i'' \sim x_i''$$

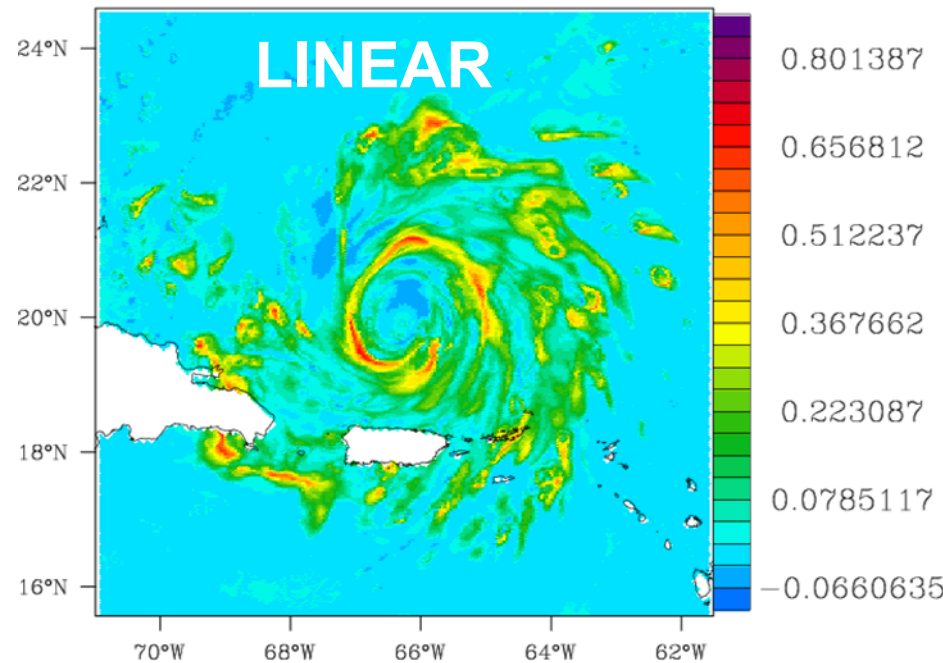
So let's try an assimilation using this observation operator:
(start with 1D-var – not quite a “retrieval”, because of the covariances)

Having started with a horizontally uniform background, each variable having the global mean value at that height level:

Avg qsnow levels 0-41, truth (g/kg)



Avg qsnow levels 0-41, anlys (g/kg)



SNOW



Second part of step 3: use nonlinear expression

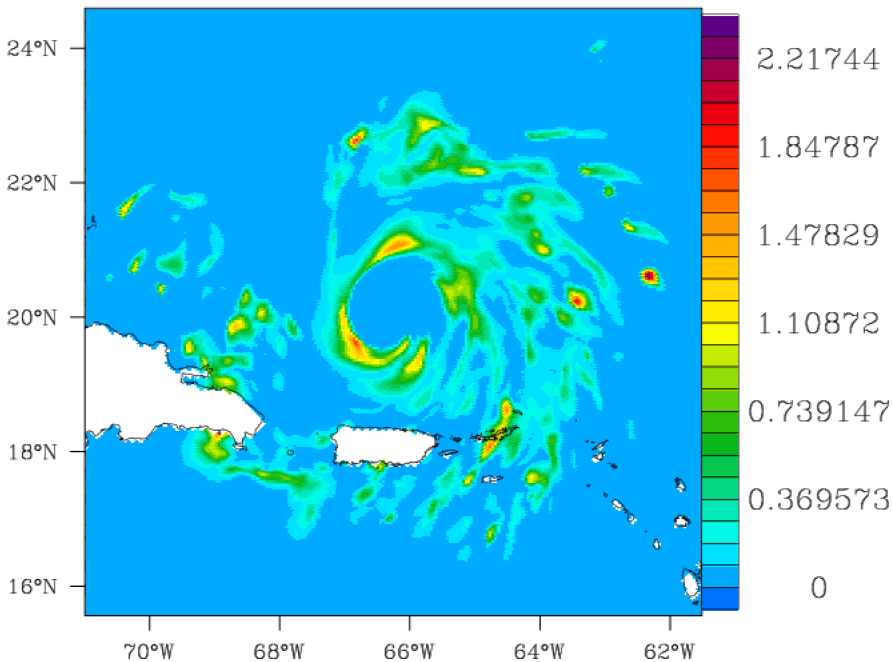
$$T_i''(x_1'', x_2'', x_3'') = \sum T_i''^{(n)} \exp(-[x_1''-x_1''^{(n)}]^2 -[x_2''-x_2''^{(n)}]^2 -[x_3''-x_3''^{(n)}]^2)$$

So let's try an assimilation using this observation operator:

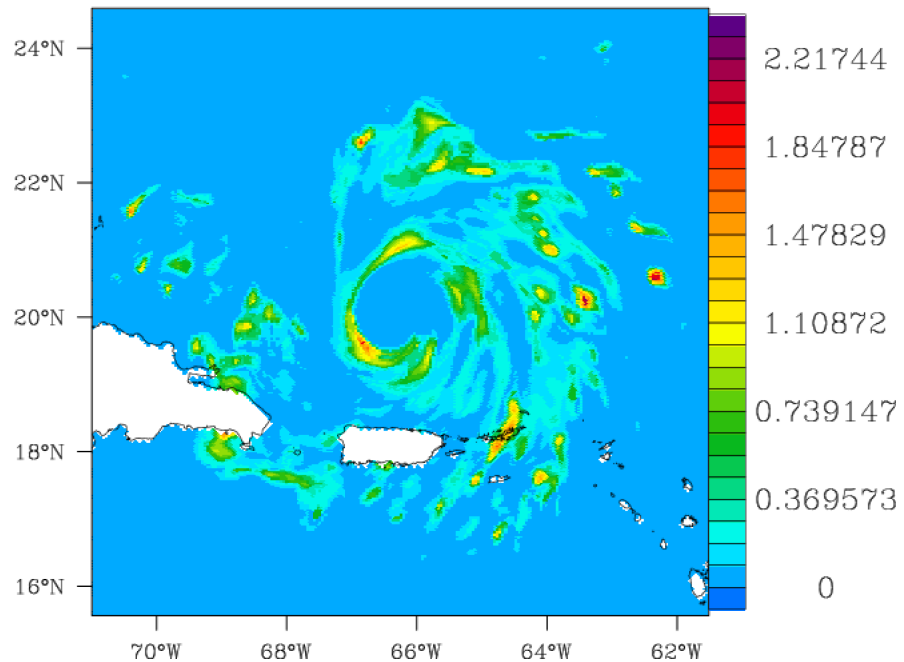
(start with 1D-var – not quite a “retrieval”, because of the covariances)

Having started with a horizontally uniform background, each variable having the global mean value at that height level:

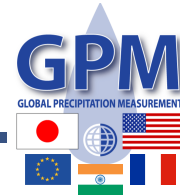
Avg totalice levels 0-41, truth (g/kg)



Avg totalice levels 0-41, anlys (g/kg)



TOTAL ICE



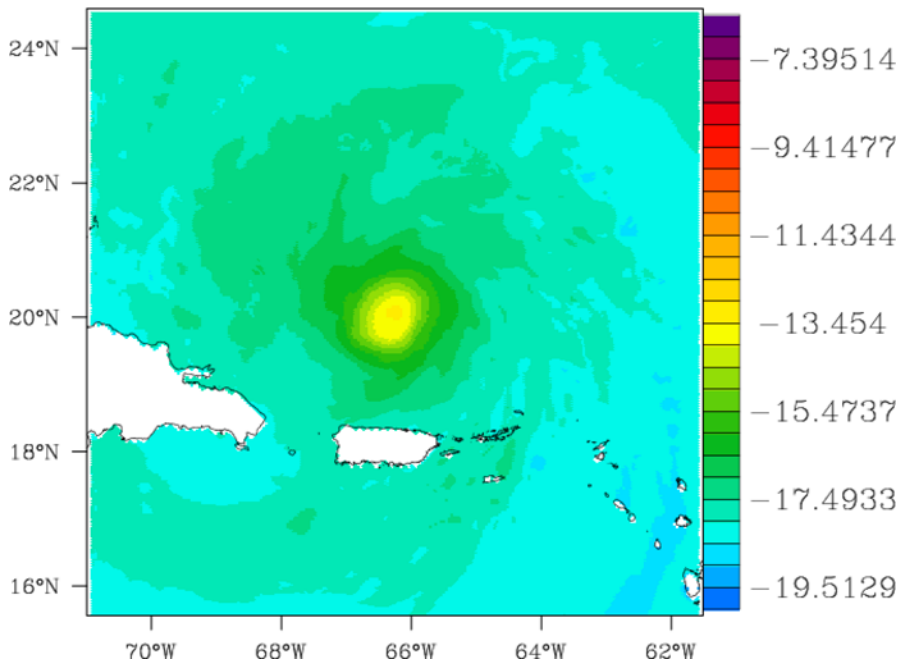
Second part of step 3: use nonlinear expression

$$T_i'' \sim x_i''$$

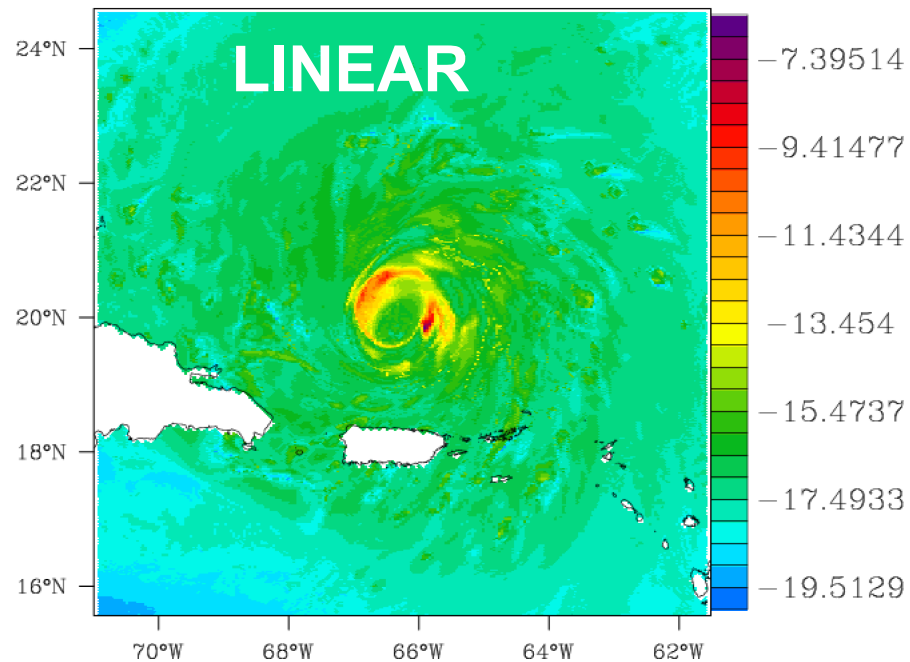
So let's try an assimilation using this observation operator:
(start with 1D-var – not quite a “retrieval”, because of the covariances)

Having started with a horizontally uniform background, each variable having the global mean value at that height level:

Avg temp levels 0-41, truth (C)



Avg temp levels 0-41, anlys (C)



temperature



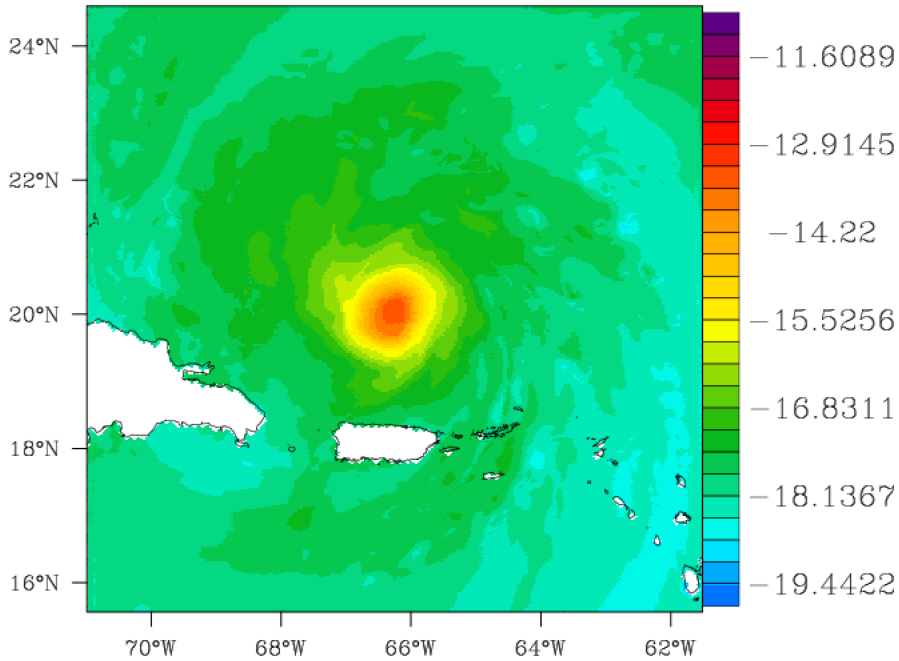
Second part of step 3: use nonlinear expression

$$T_i''(x_1'', x_2'', x_3'') = \sum T_i''^{(n)} \exp(-[x_1''-x_1''^{(n)}]^2 -[x_2''-x_2''^{(n)}]^2 -[x_3''-x_3''^{(n)}]^2)$$

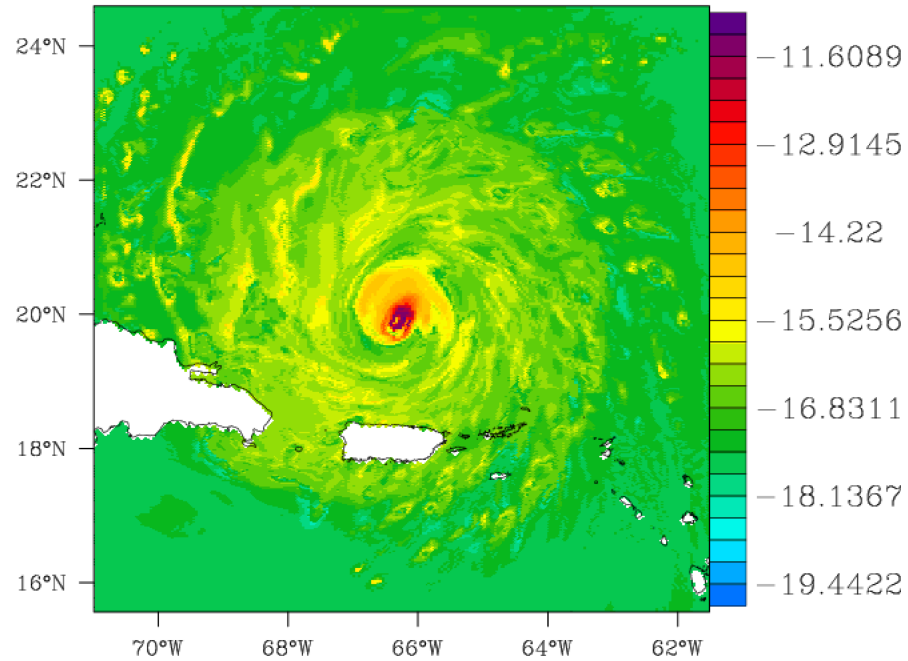
So let's try an assimilation using this observation operator:
(start with 1D-var – not quite a “retrieval”, because of the covariances)

Having started with a horizontally uniform background, each variable having the global mean value at that height level:

Avg temp levels 0-41, truth (C)



Avg temp levels 0-41, anlys (C)



Look P/Ma! a warm core! 45

temperature



Cast:

Z.S. Haddad (JPL)

Jeff Steward (UCLA)

Joe Munchak (GSFC)

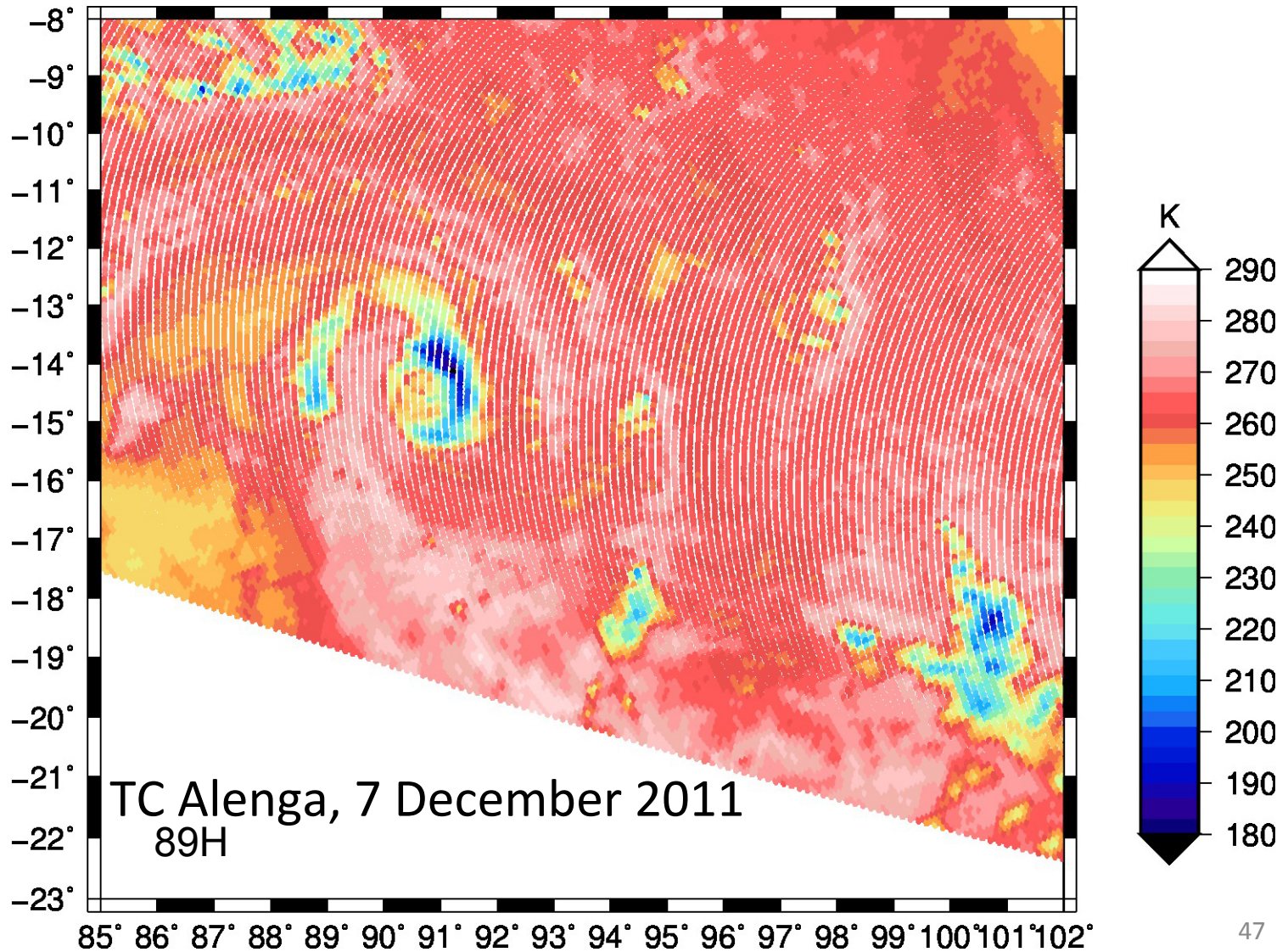
Sahra Kacimi (JPL)

Hsiao-Chieh Tseng (UC Davis)

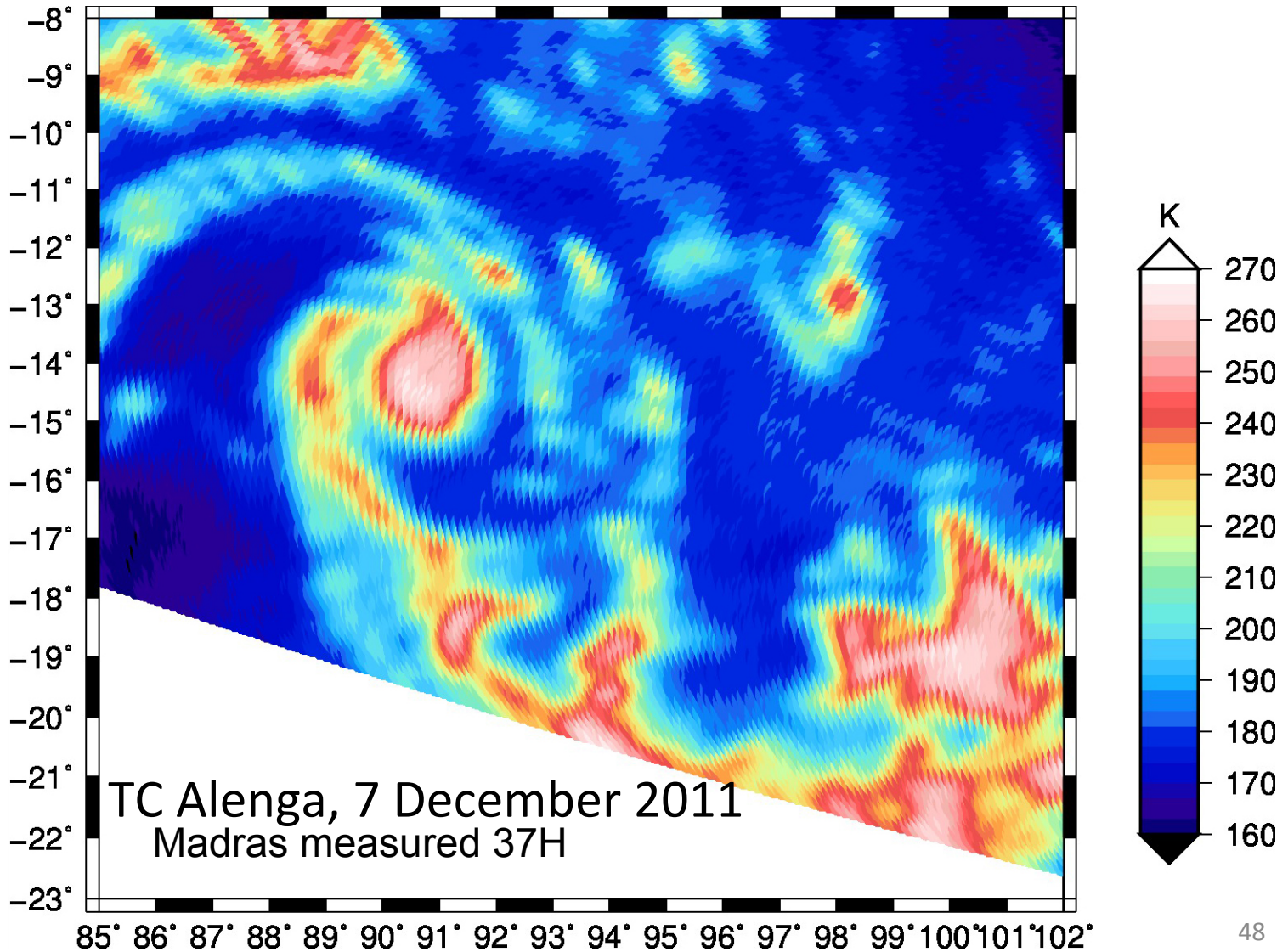
Svetla Hristova-Veleva (JPL)

Shu-Hua Chen (UC Davis)

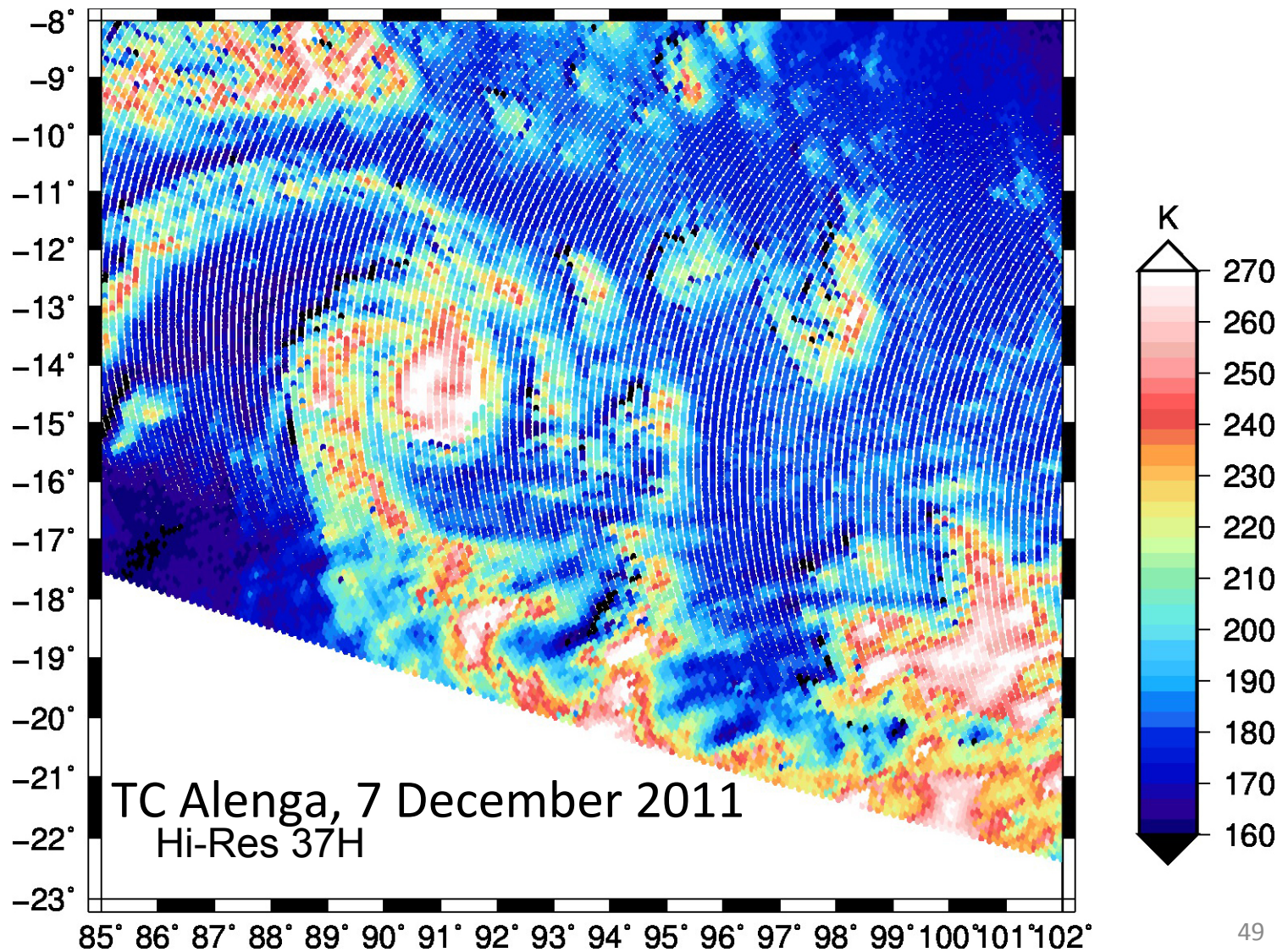
$$T_{hiRes} = t(89_{hiRes}) + (1 + C P^t E^{-1} P)^{-1} C P^t E^{-1} [T_{loRes} - P t(89_{hiRes})]$$



$$T_{hiRes} = t(89_{hiRes}) + (1 + C P^t E^{-1} P)^{-1} C P^t E^{-1} [T_{loRes} - P t(89_{hiRes})]$$



$$T_{hiRes} = t(89_{hiRes}) + (1 + C P^t E^{-1} P)^{-1} C P^t E^{-1} [T_{loRes} - P t(89_{hiRes})]$$





$$T_{hiRes} = t(89_{hiRes}) + (1 + C P^t E^{-1} P)^{-1} C P^t E^{-1} [T_{loRes} - P t(89_{hiRes})]$$

We now need to analyze the dependence on

1. the pixel classification conditioned on T_{89V} , T_{89V} ,
2. the derived mean relations (t) between (T_{89V} , T_{89V}) & the lower-frequency brightness temperatures (at same resolution),
3. the quantification of the covariance around (C) the mean relations
4. the allowable error on the deconvolution portion (E)
5. the size of the domain



Two ways to account for the dependence of microwave observations of precipitation on uncertain microphysics

Algorithm Theoretical Basis (“Default Algorithm”)

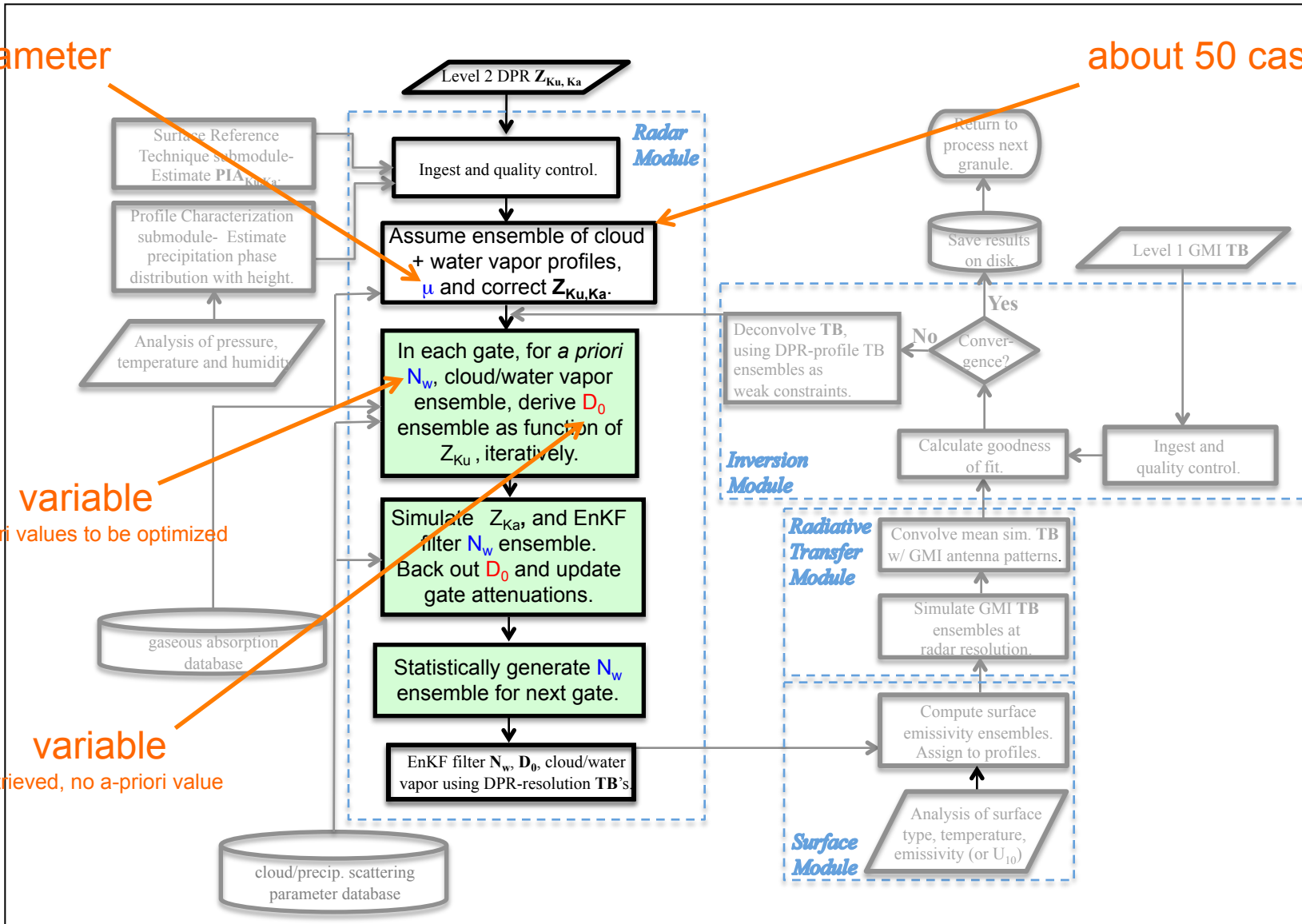


parameter

about 50 cases

variable
a-priori values to be optimized

variable
retrieved, no a-priori value





Two ways to account for the dependence of microwave observations of precipitation on uncertain microphysics

Algorithm Theoretical Basis (“Default Algorithm”)

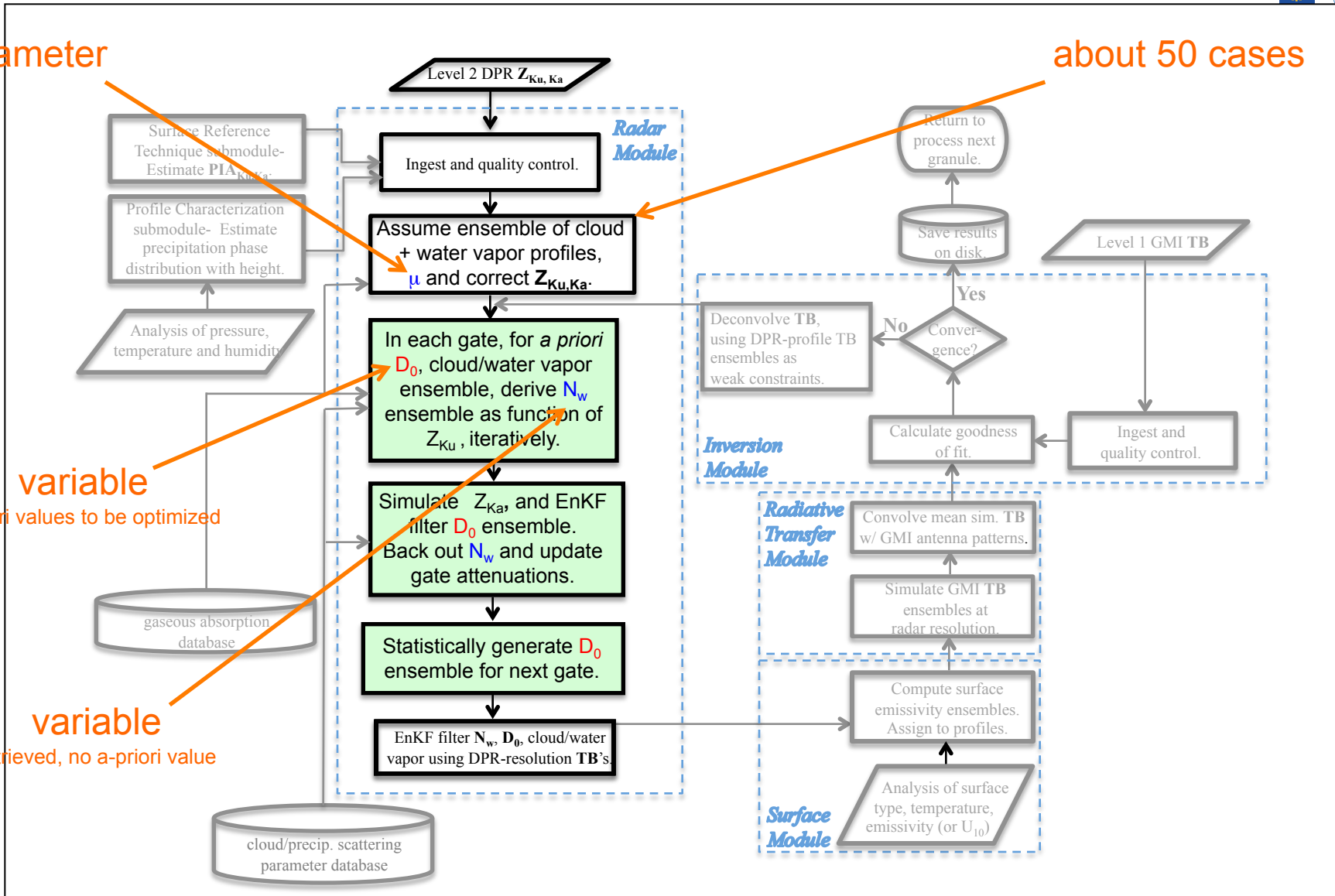


parameter

about 50 cases

variable
a-priori values to be optimized

variable
retrieved, no a-priori value





Test alternative



about 50 cases

parameter

variable
a-priori values to be optimized

variable
retrieved, no a-priori value

